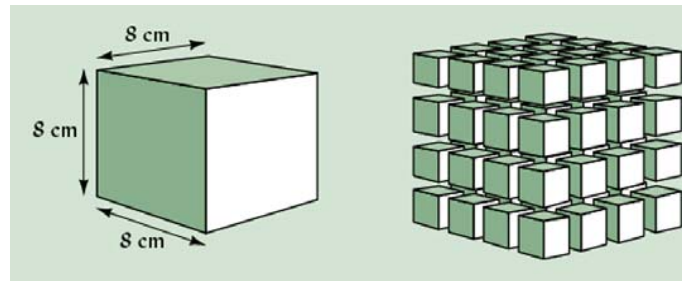


# The Colloid Fraction of Soil

- Size – 1  $\mu\text{m}$  in size (2  $\mu\text{m}$ ?)

	0.002		0.05	0.10	0.25	0.5	1.0	2.0	
Clay	Silt				Very fine	Fine	Med.	Coarse	Very coarse
			Sand						

- Surface Area



- Surface Charge (generally negative)
- Cation (and Anion) Adsorption

International  
Society of  
Soil Science

Clay	Silt	Sand				Gravel
		Fine		Coarse		
0.002	0.02	0.2		2.0		
0.002	0.05	0.10	0.25	0.5	1.0	
				2.0		

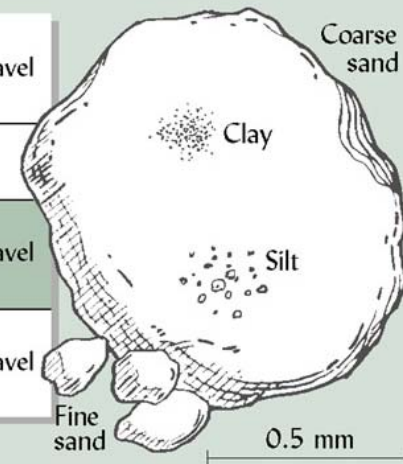
United States  
Department of  
Agriculture

Clay	Silt	Very fine	Fine	Med.	Coarse	Very coarse	Gravel
		Sand					

United States  
Public Roads  
Administration

Clay	Silt	Sand			Gravel
		Fine		Coarse	
0.005	0.05	0.25		2.0	

Particle diameter (mm, log scale)



RELATIVE AMOUNT

**primary  
minerals**

quartz, feldspars,  
micas, etc.

**secondary  
minerals**

layer silicate clays,  
oxides,  
non-crystalline  
aluminosilicates

SAND

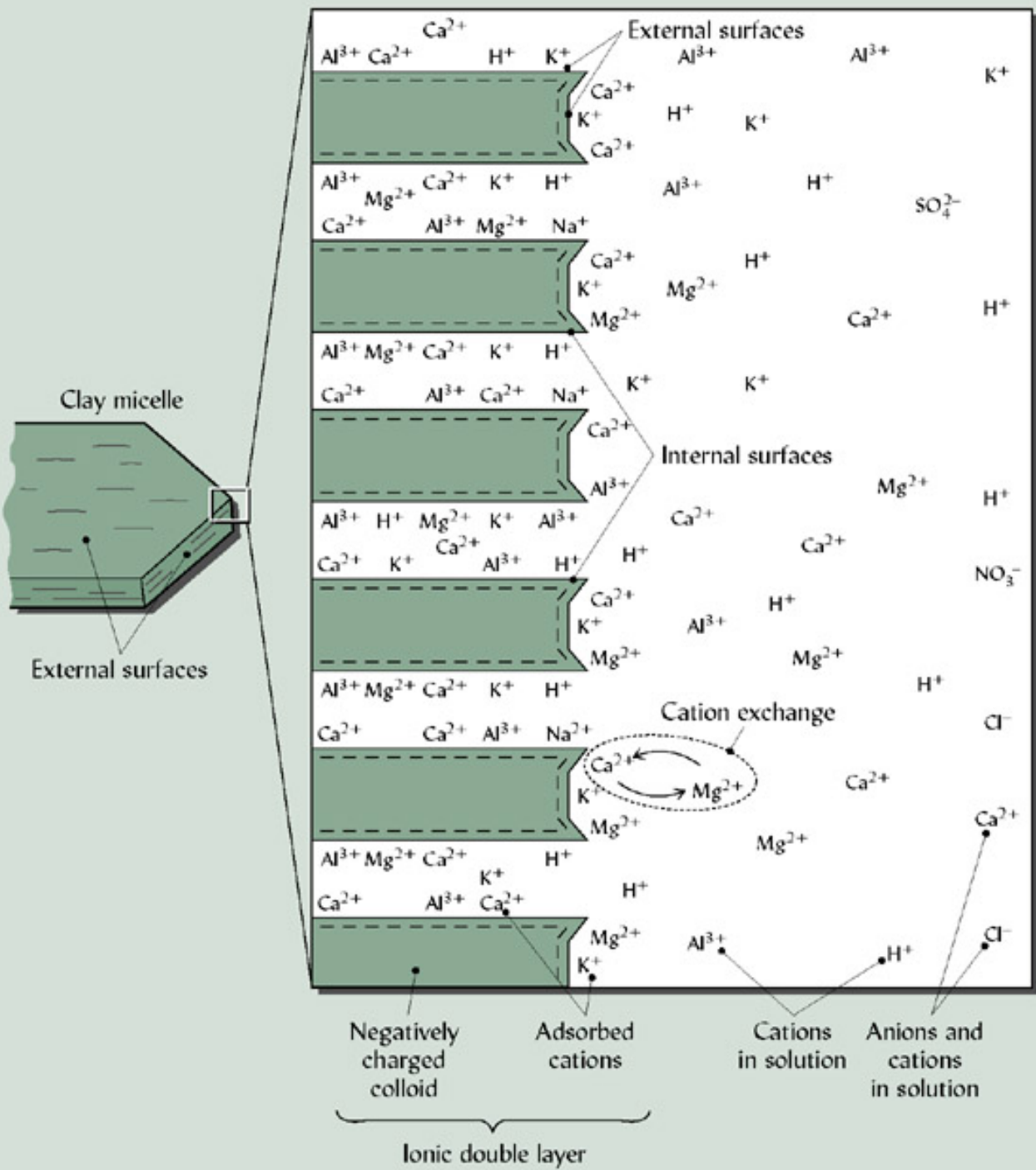
50

SILT

2

CLAY

PARTICLE SIZE ( $\mu\text{m}$ )

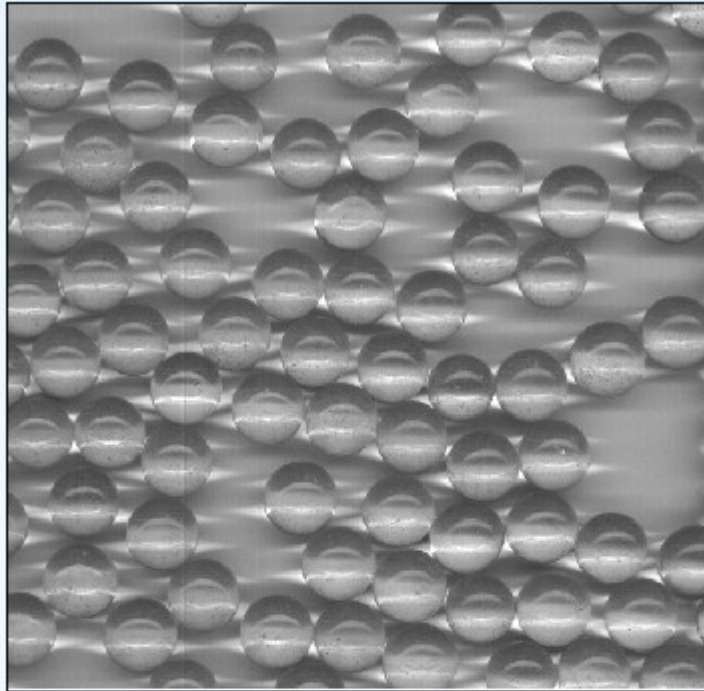


# Types of Soil Colloids

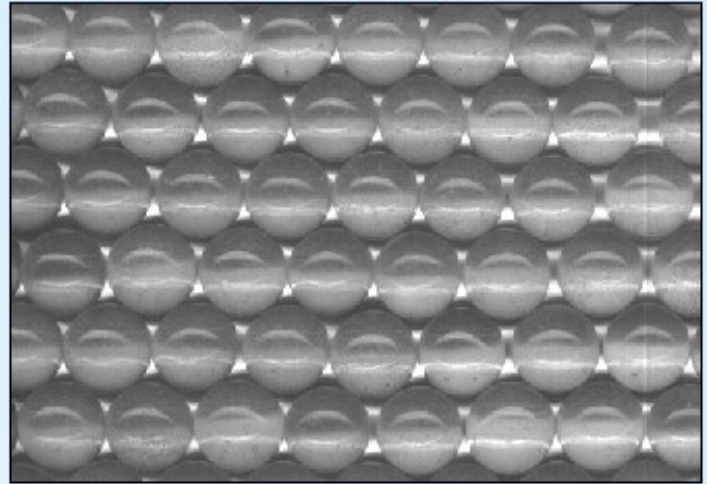
- **Crystalline silicate clays**
  - Phyllosilicates → tetrahedral and octahedral crystal sheets
- **Non-crystalline silicate clays (Andisols)**
  - Dominately amorphous clays (allophane and imogolite)
- **Iron and aluminum oxides (Oxisols & ...)**
  - Dominately gibbsite (Al-oxide) and goethite (Fe-oxide)
- **Organic (humus) colloids (Histosols &...)**
  - Non-crystalline colloids dominated by long C-chain molecules

# Types of Soil Colloids

- **Crystalline silicate clays**
  - Phyllosilicates → tetrahedral and octahedral crystal sheets
- Non-crystalline silicate clays (Andisols)
  - Dominately amorphous clays (allophane and imogolite)
- Iron and aluminum oxides (Oxisols & ...)
  - Dominately gibbsite (Al-oxide) and goethite (Fe-oxide)
- Organic (humus) colloids (Histosols &...)
  - Non-crystalline colloids dominated by long C-chain molecules



- Ions dissolved in magma or water move rapidly
- No regular pattern



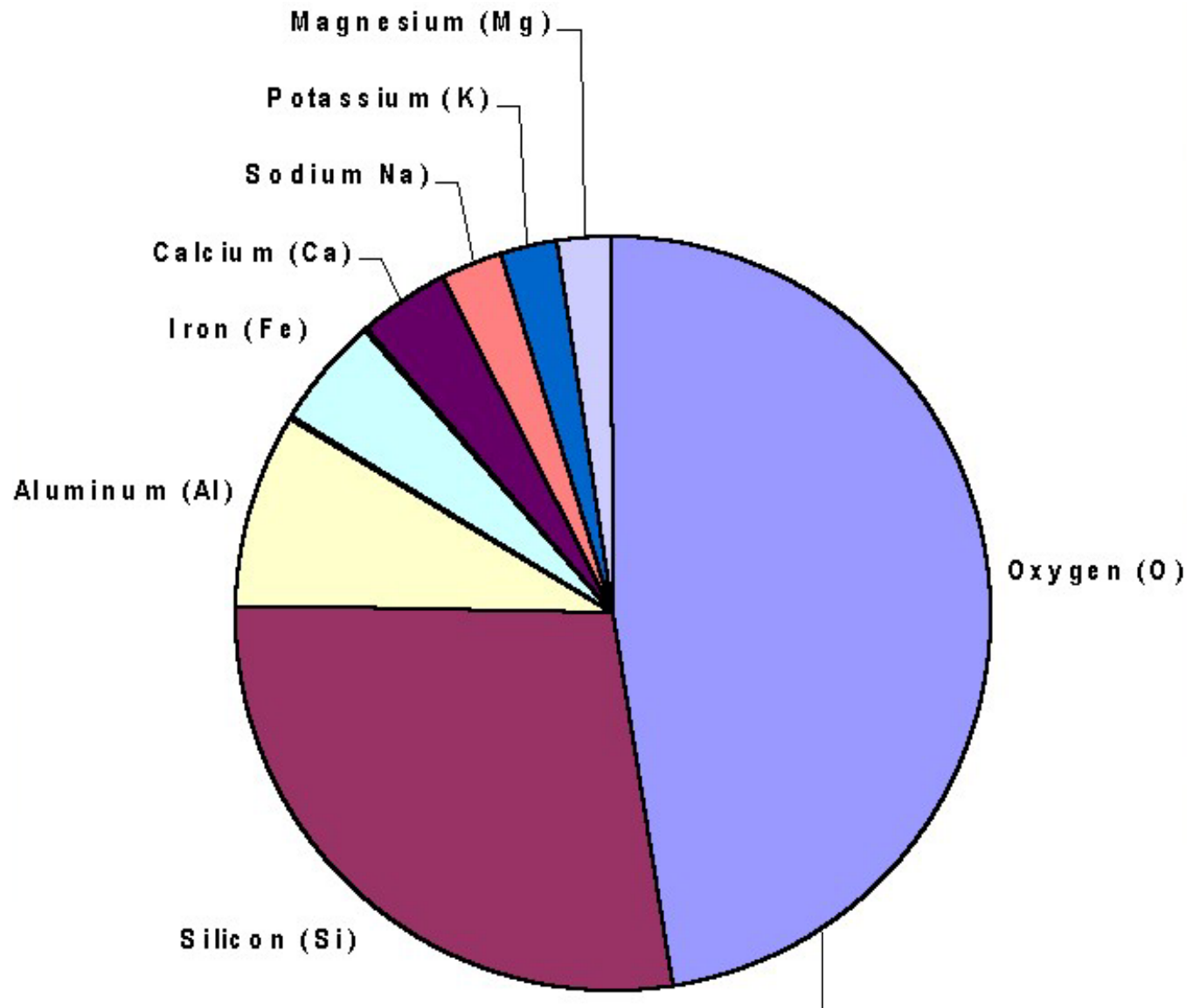
- Ions “frozen” in solid crystal have symmetrical repeating pattern

# Chemical Weathering

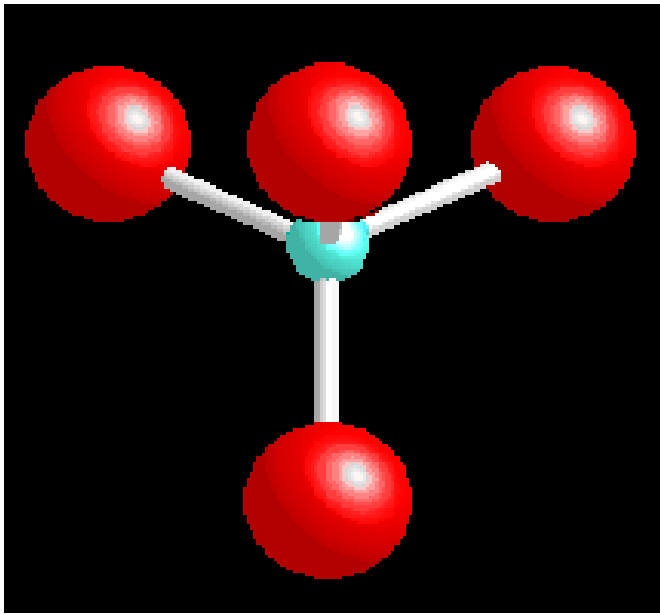
Process that changes minerals from their original composition to a new composition by:

- Hydrolysis – addition of a  $\text{H}^+$  to the structure
- Hydration – addition of a water molecule
- Oxidation / Reduction – gain or loss of an electron
- Dissolution / Carbonation –  $\text{H}^+$  from  $\text{H}_2\text{CO}_3$

# Composition of the Earth's Crust

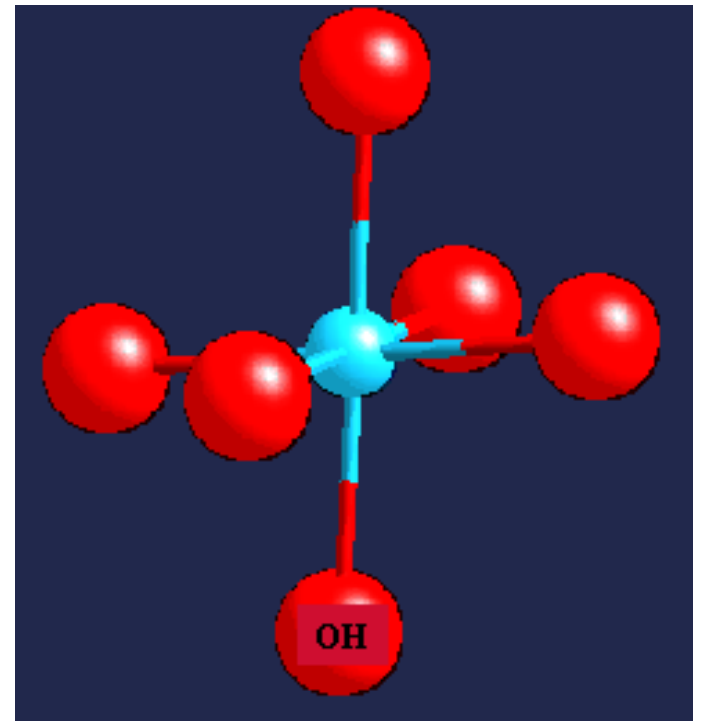


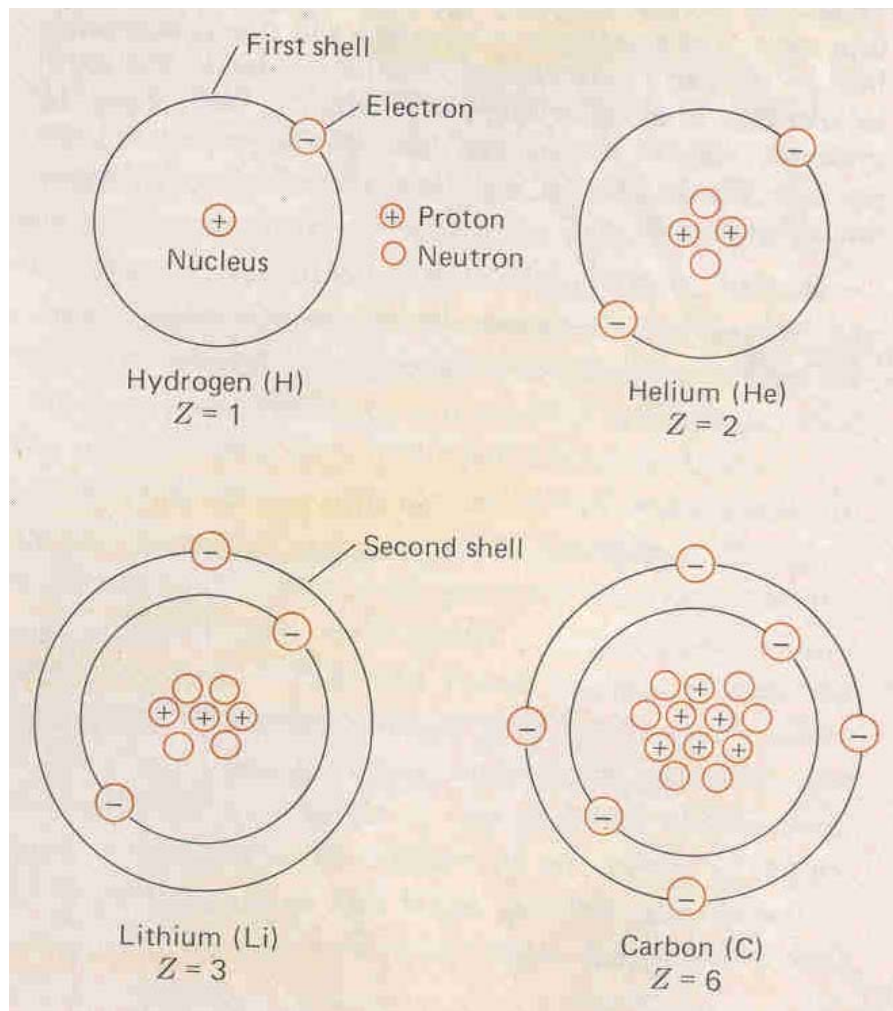




Tetrahedra - Si

Octahedra – Al or Mg

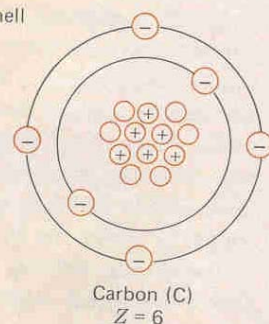
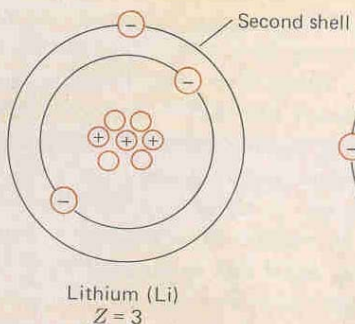
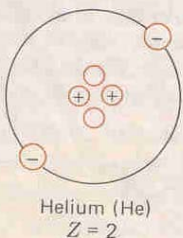
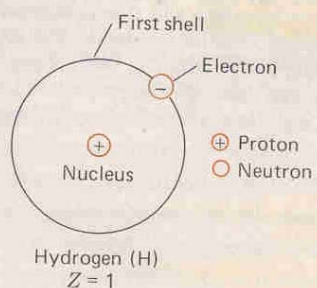




Energy level	# of electrons needed to fill shell
1	2
2	8
3	8
4	8
5	8

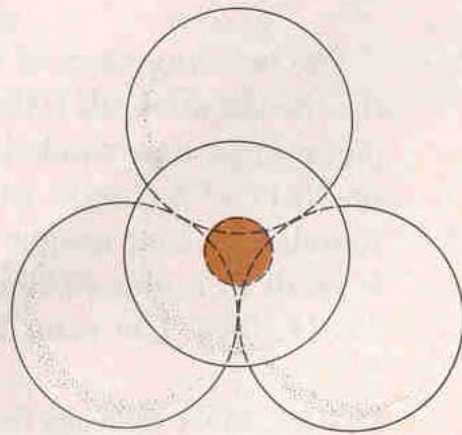
Number of electrons in outermost shell

Energy level	Electron shell	1							2
		1							2
1	K	1 H Hydrogen							2 He Helium
		1	2	3	4	5	6	7	8
2	L	3 Li Lithium	4 Be Beryllium	5 B Boron	6 C Carbon	7 N Nitrogen	8 O Oxygen	9 F Fluorine	10 Ne Neon
3	M	11 Na Sodium	12 Mg Magnesium	13 Al Aluminum	14 Si Silicon	15 P Phosphorus	16 S Sulfur	17 Cl Chlorine	18 Ar Argon
4	N	19 K Potassium	20 Ca Calcium	31 Ga Gallium	32 Ge Germanium	33 As Arsenic	34 Se Selenium	35 Br Bromine	36 Kr Krypton
		29 Cu Copper	26 Fe Iron						
Covalence		1	2	3	4	3	2	1	Inert; octet filled
		Lose electrons to leave an octet in next lower shell			Gain electrons to complete an octet				

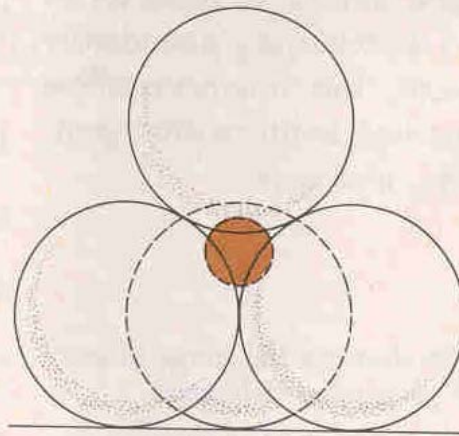


Energy level      # of electrons  
needed to fill shell

1	2
2	8
3	8
4	8
5	8

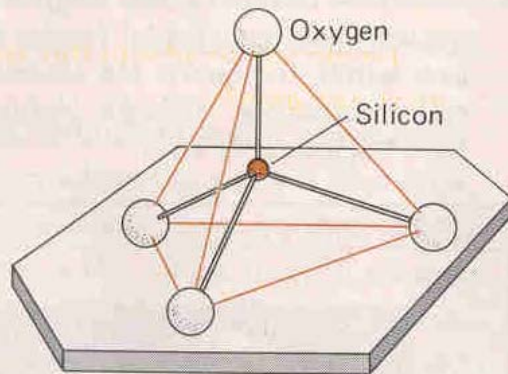


Top view

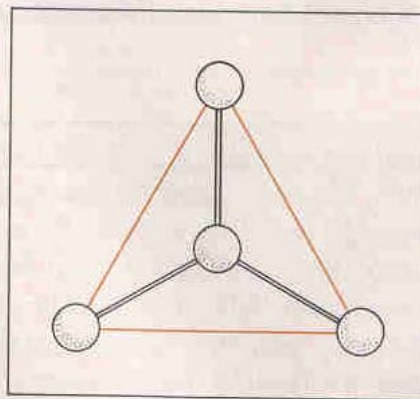


Side view

A. Tetrahedron, ions to scale, close packing

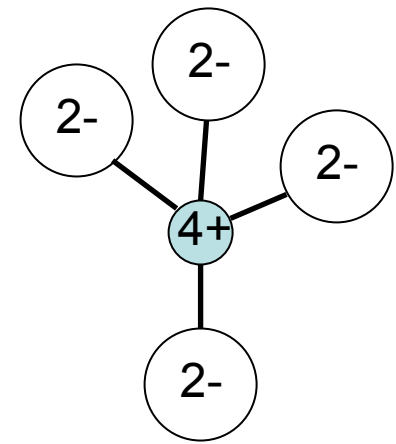


Perspective view from side

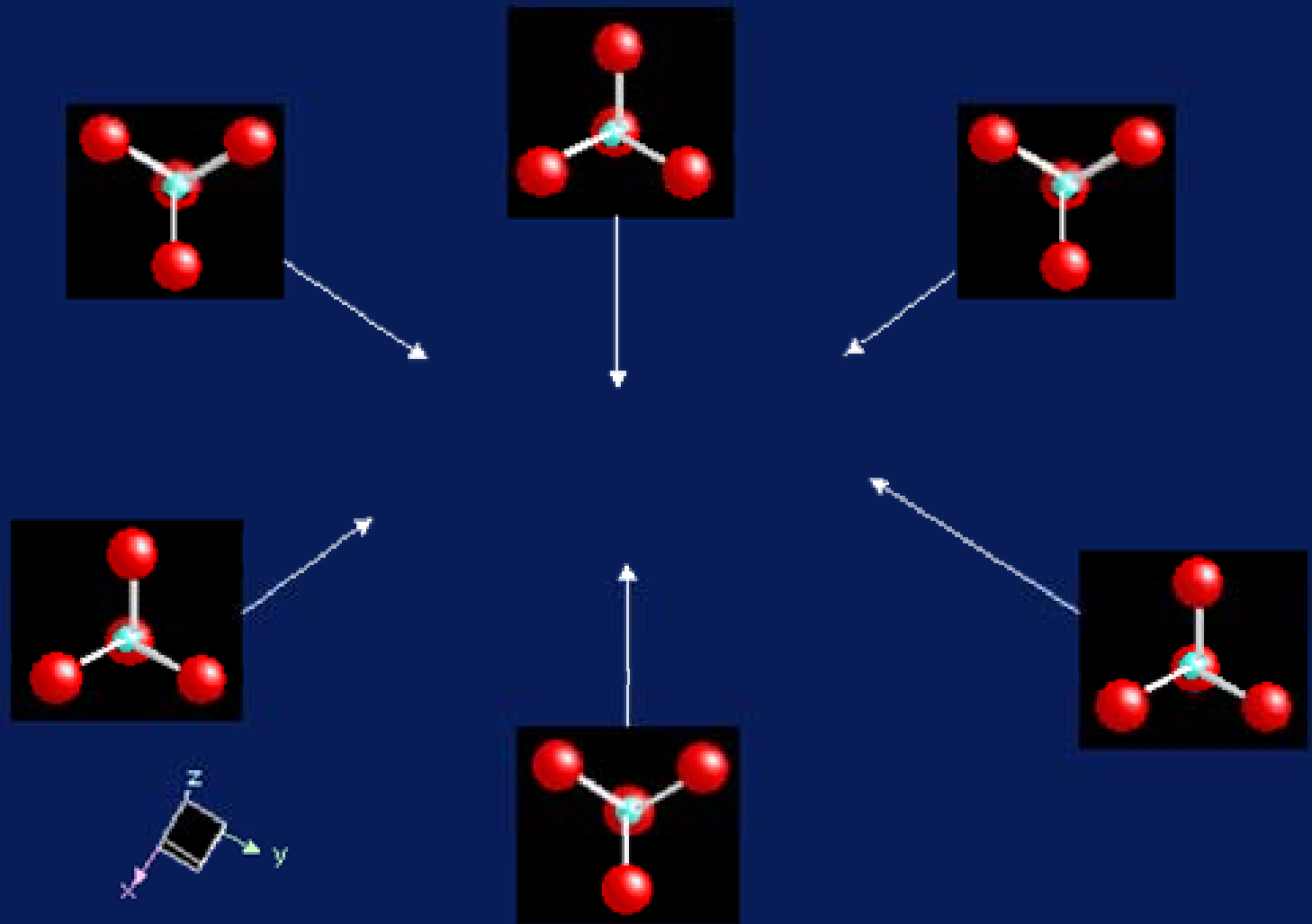


Looking straight down

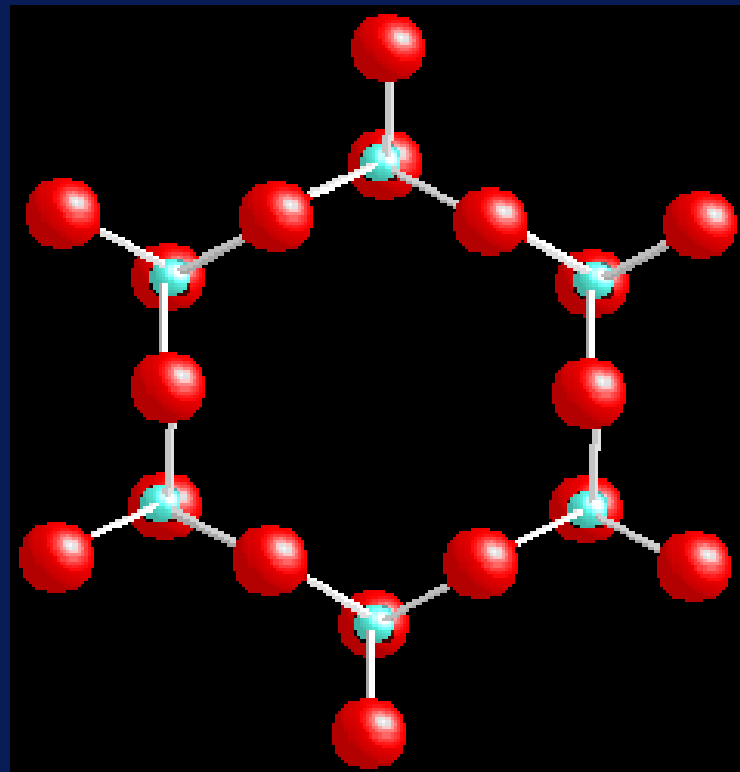
B. Expanded tetrahedron



# CONNECT THE TETRAHEDRA

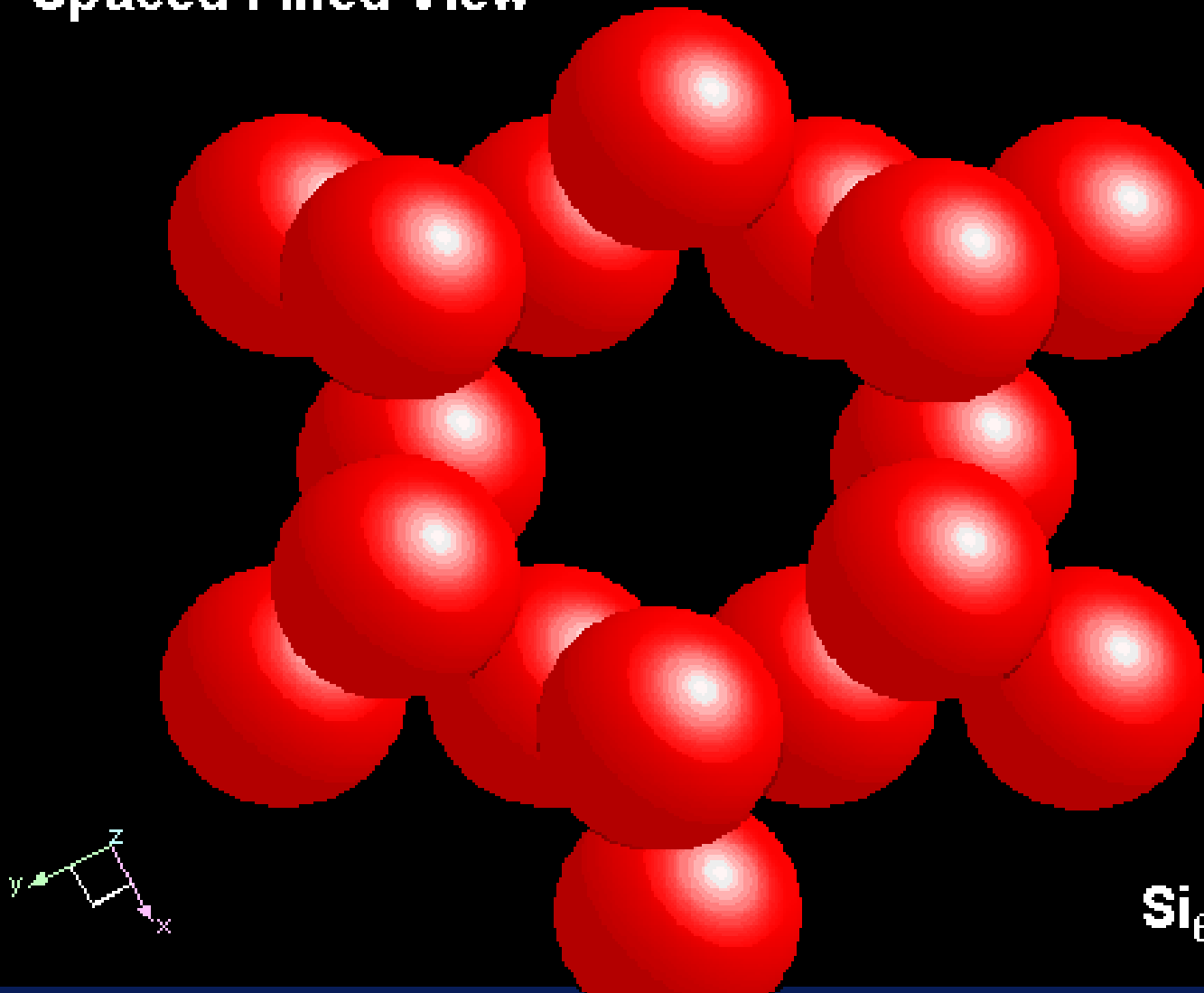


# SILICA TETRAHEDRAL RING

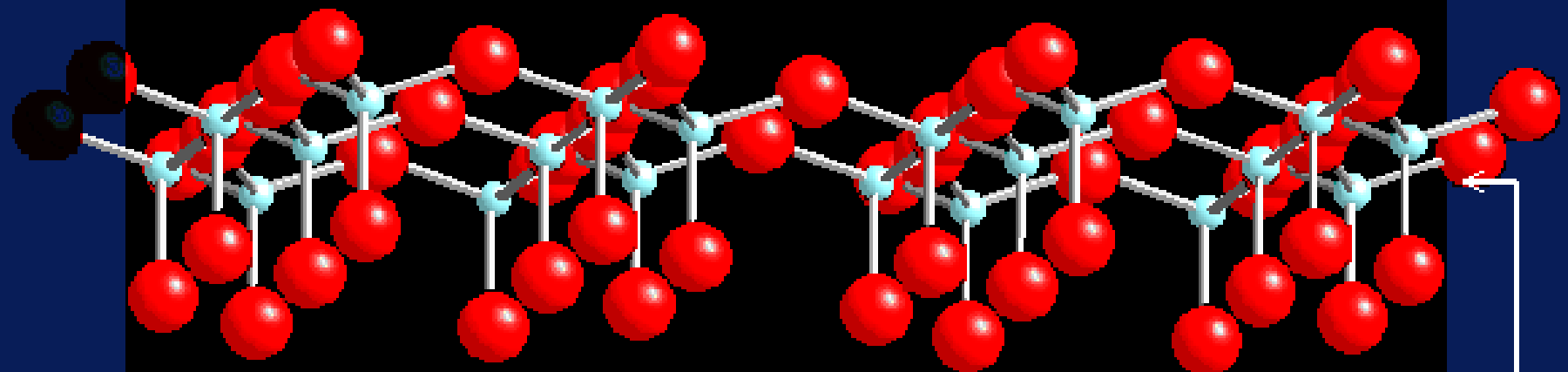


# SILICA TETRAHEDRAL RING

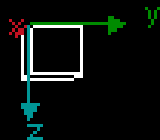
Spaced Filled View



# TETRAHEDRAL SHEET

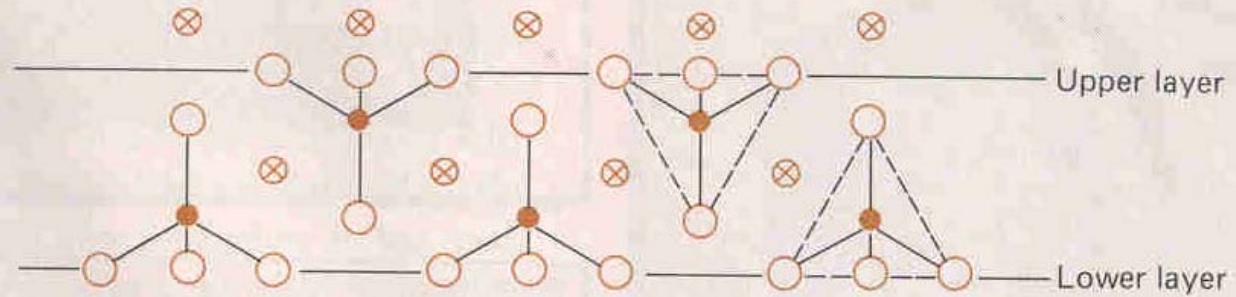


unsaturated apical  
oxygen

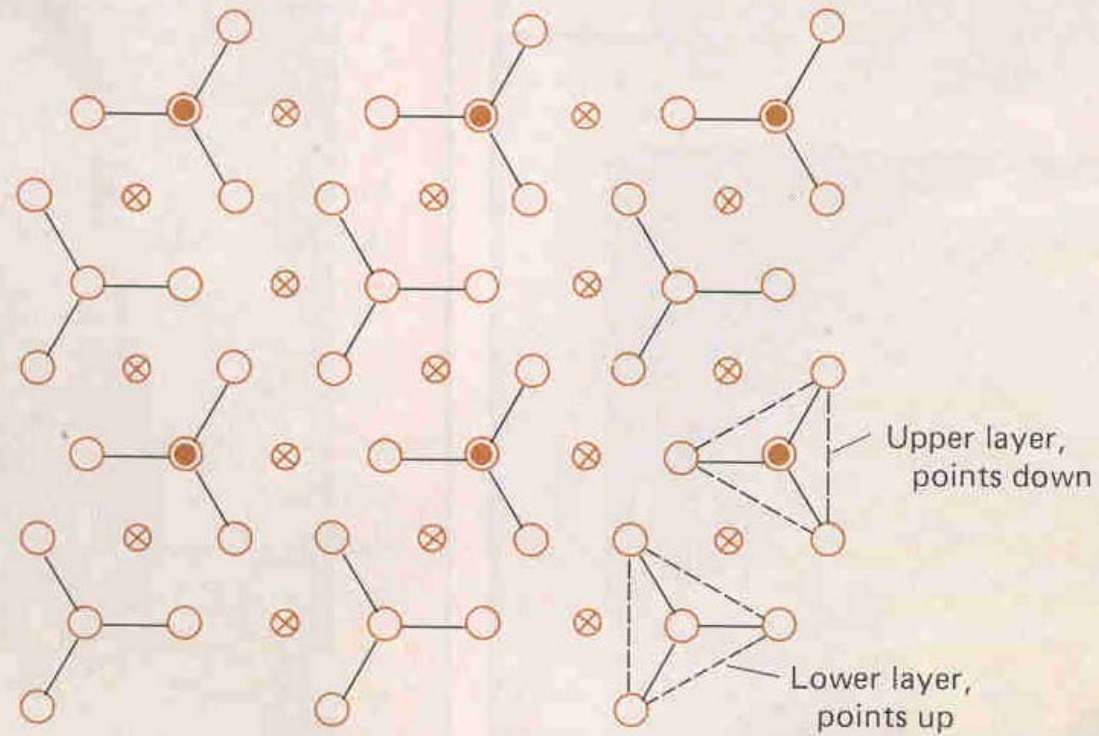




Side view:



Top view, looking down on layers:

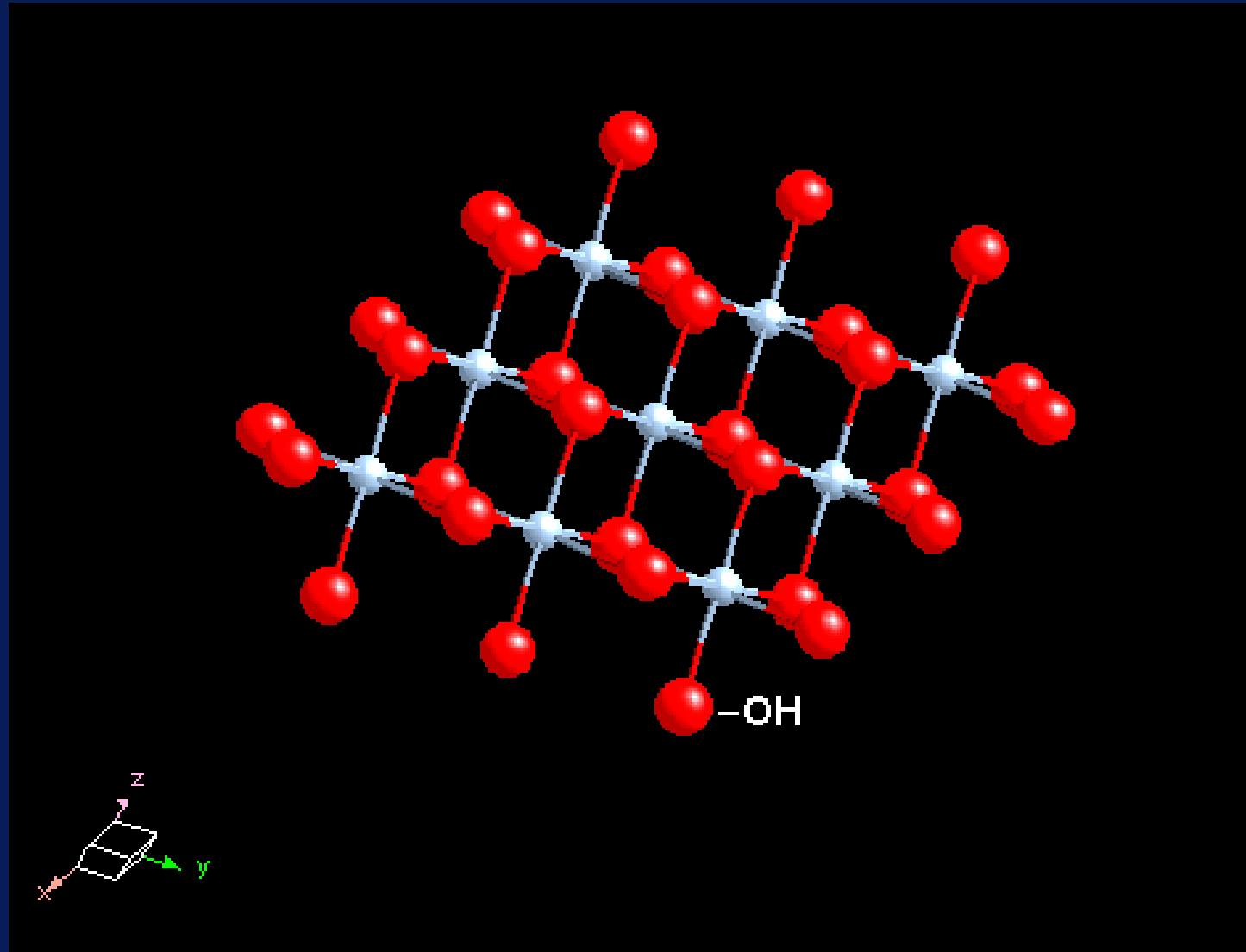


○ Oxygen

● Silicon

⊗ Magnesium, iron

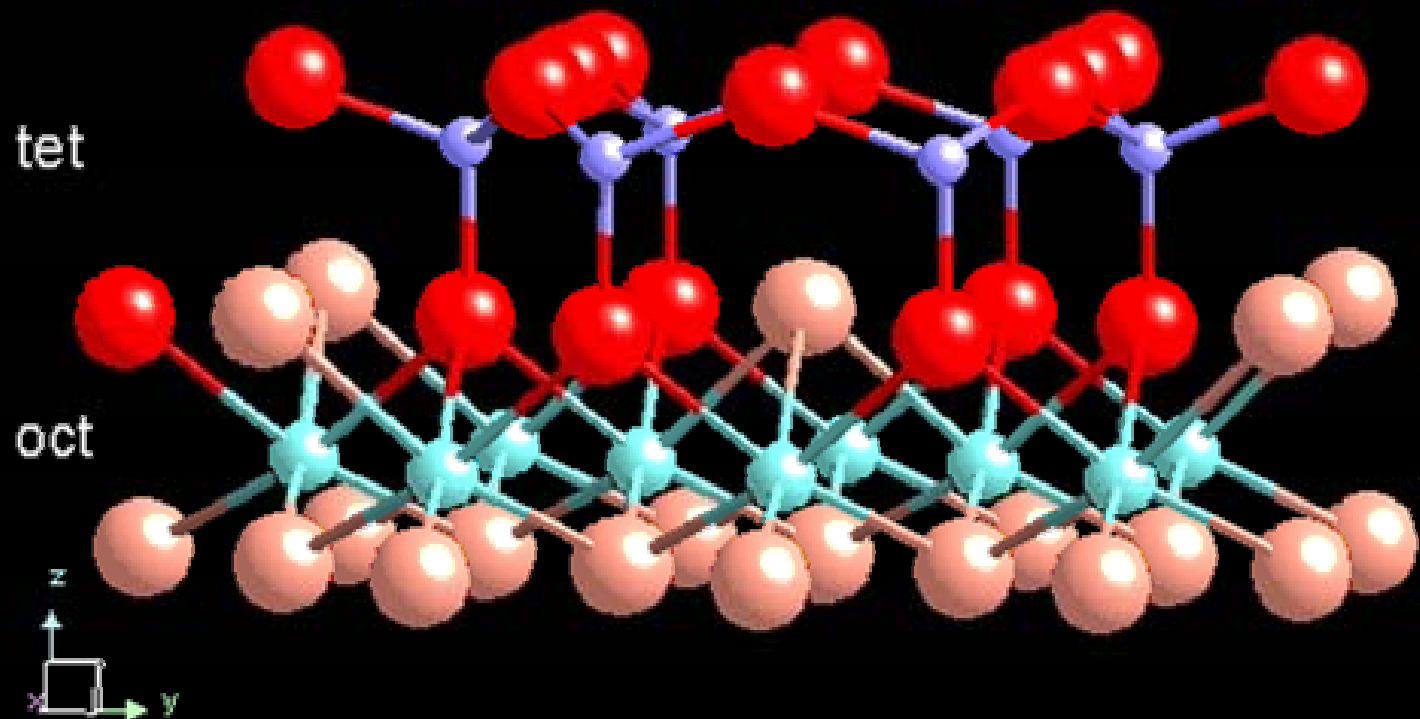
# OCTAHEDRAL SHEET



# OCTAHEDRAL-TETRAHEDRAL LINKAGE

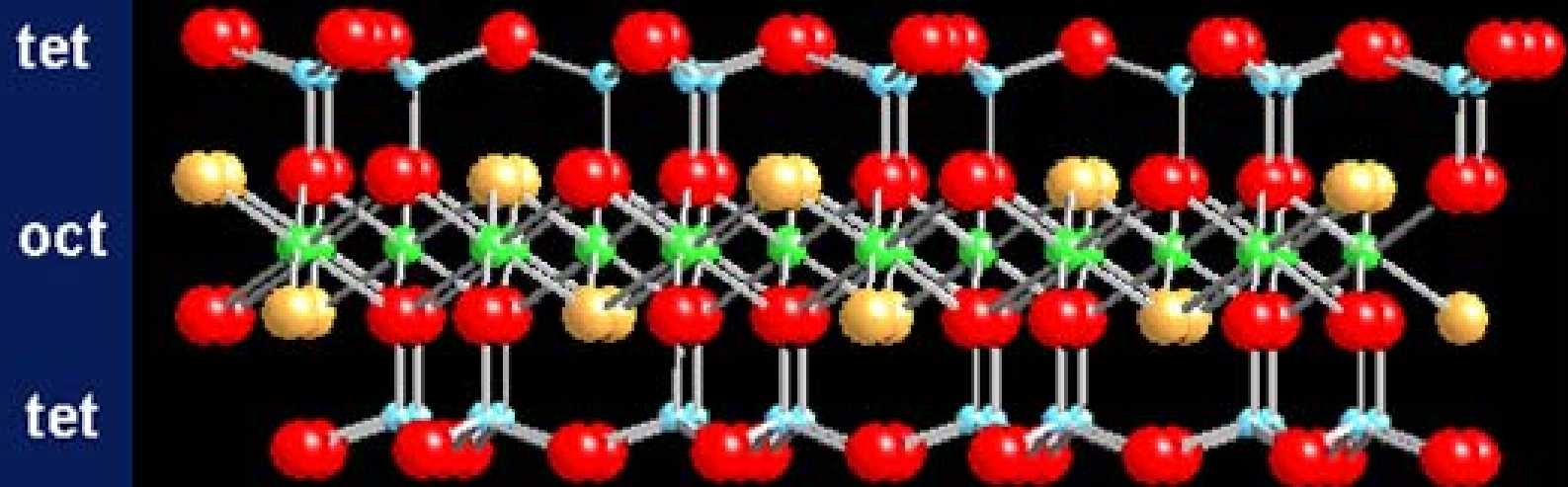
Sharing of Apical Oxygens in Tetrahedral Sheet  
with Hydroxyls of Octahedral Sheet

Serpentine (1:1 trioctahedral mineral)

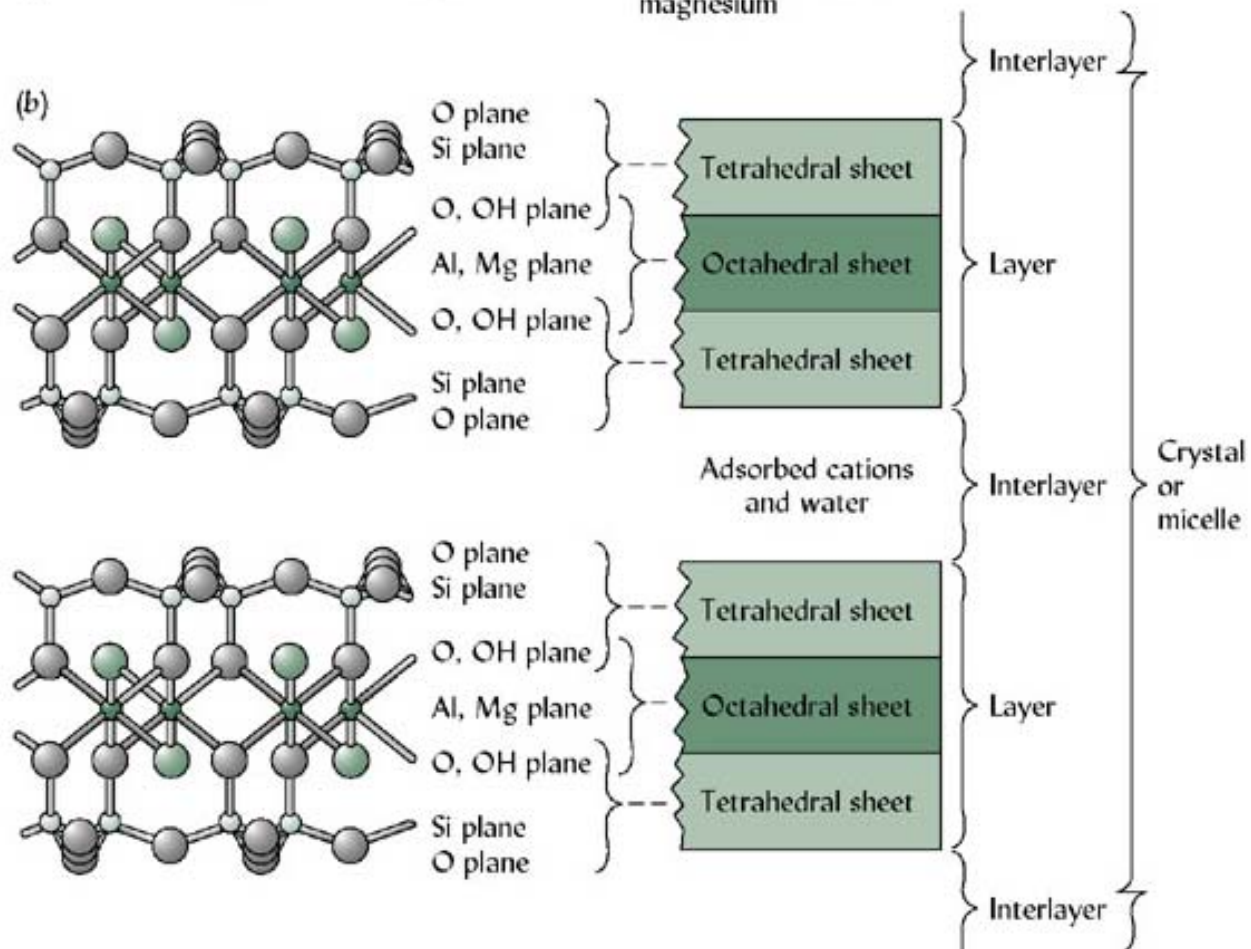
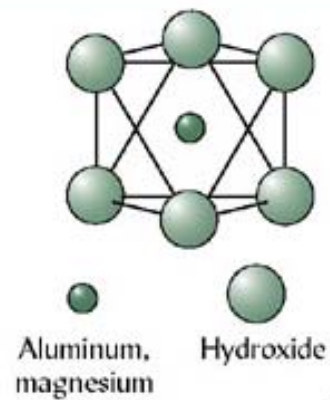
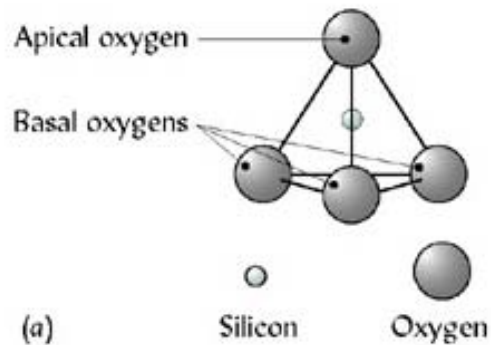


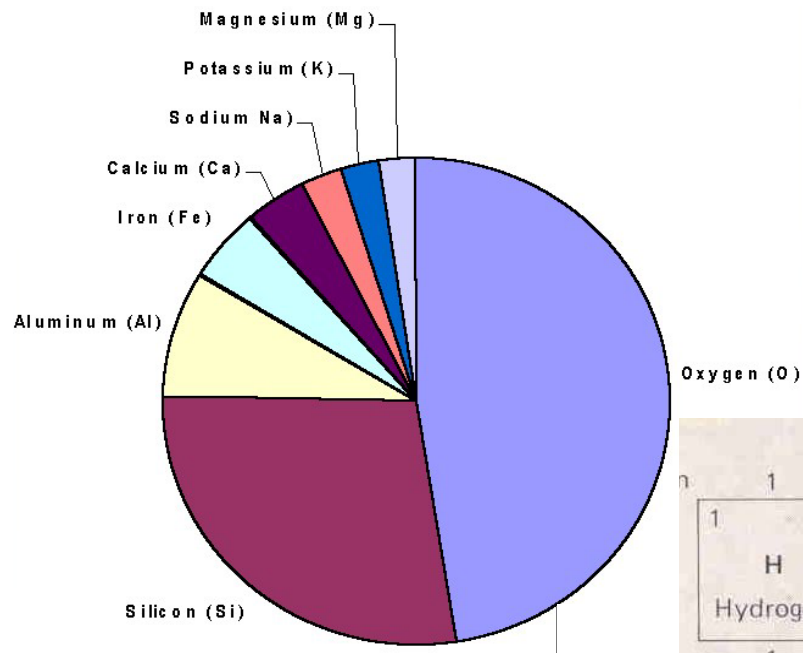
# OCTAHEDRAL-TETRAHEDRAL LINKAGE

Sharing of Apical Oxygens in Tetrahedral Sheet with Hydroxyls of Two Octahedral Sheets



Talc (2:1 trioctahedral mineral)

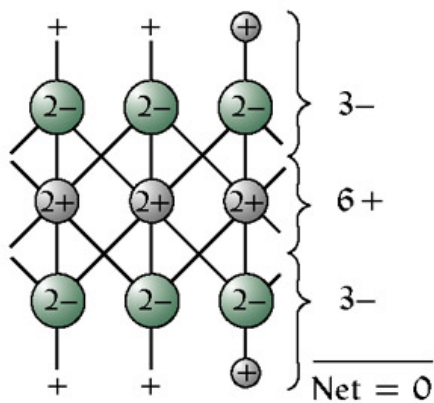




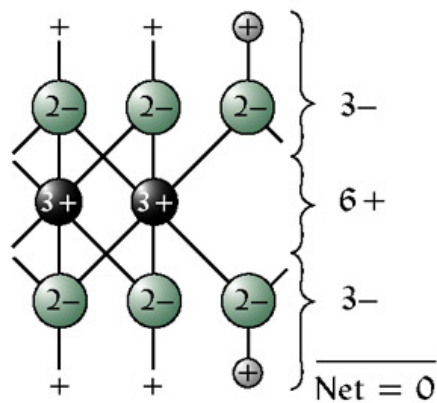
Number of electrons in outermost shell

1		2					
Atomic number		Atomic number					
1	H Hydrogen	2 He Helium					
1	2	3	4	5	6	7	8
3	4	5	6	7	8	9	10
Li Lithium	Be Beryllium	B Boron	C Carbon	N Nitrogen	O Oxygen	F Fluorine	Ne Neon
11	12	13	14	15	16	17	18
Na Sodium	Mg Magnesium	Al Aluminum	Si Silicon	P Phosphorus	S Sulfur	Cl Chlorine	Ar Argon
19	20	31	32	33	34	35	36
K Potassium	Ca Calcium	Ga Gallium	Ge Germanium	As Arsenic	Se Selenium	Br Bromine	Kr Krypton
29	26						
Cu Copper	Fe Iron						
1	2	3	4	3	2	1	Inert; octet filled
Lose electrons to leave an octet in next lower shell				Gain electrons to complete an octet			

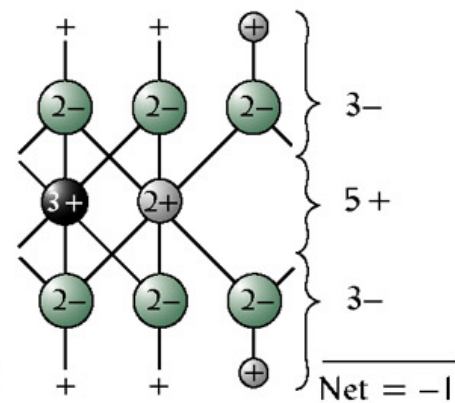




Trioctahedral  
(3 cations)



Dioctahedral  
(2 cations)



Dioctahedral  
with isomorphic  
substitution



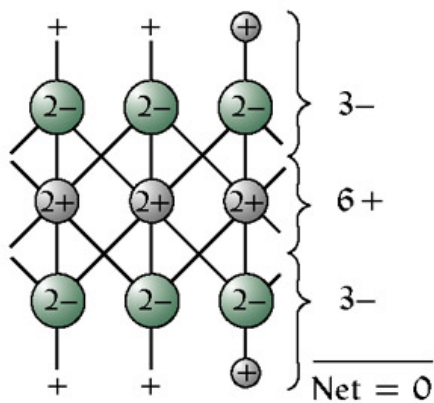
**TABLE 8.2** Ionic Radii of Elements Found in Silicate Clays and an Indication of Which Are Found in the Tetrahedral and Octahedral Sheets

*Note that Al, Fe, O, and OH can fit in either.*

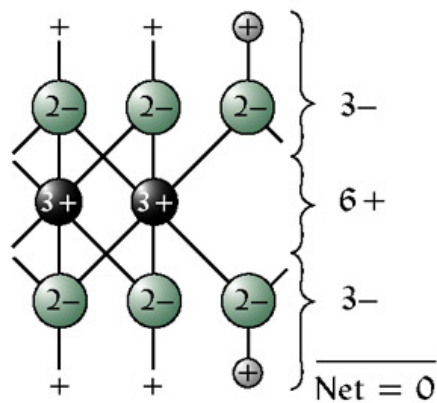
<i>Ion</i>	<i>Radius, nm<sup>a</sup></i>	<i>Found in</i>
Si <sup>4+</sup>	0.042	Tetrahedral sheet
Al <sup>3+</sup>	0.051	
Fe <sup>3+</sup>	0.064	
Mg <sup>2+</sup>	0.066	Octahedral sheet
Zn <sup>2+</sup>	0.074	
Fe <sup>2+</sup>	0.070	Exchange sites
Na <sup>+</sup>	0.097	
Ca <sup>2+</sup>	0.099	
K <sup>+</sup>	0.133	Both sheets
O <sup>2-</sup>	0.140	
OH <sup>-</sup>	0.155	

<sup>a</sup> 1 nm = 10<sup>-9</sup>m.

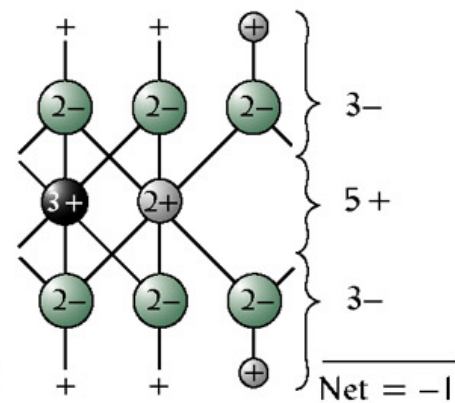




Trioctahedral  
(3 cations)



Dioctahedral  
(2 cations)

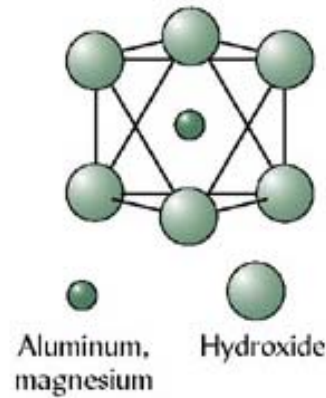
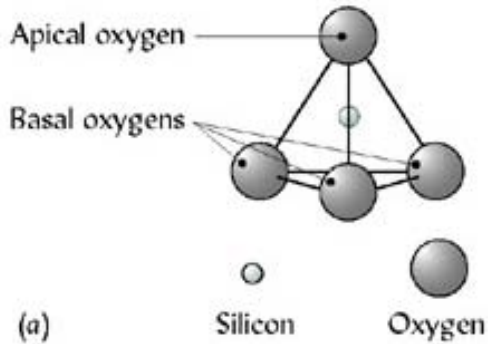


Dioctahedral  
with isomorphic  
substitution

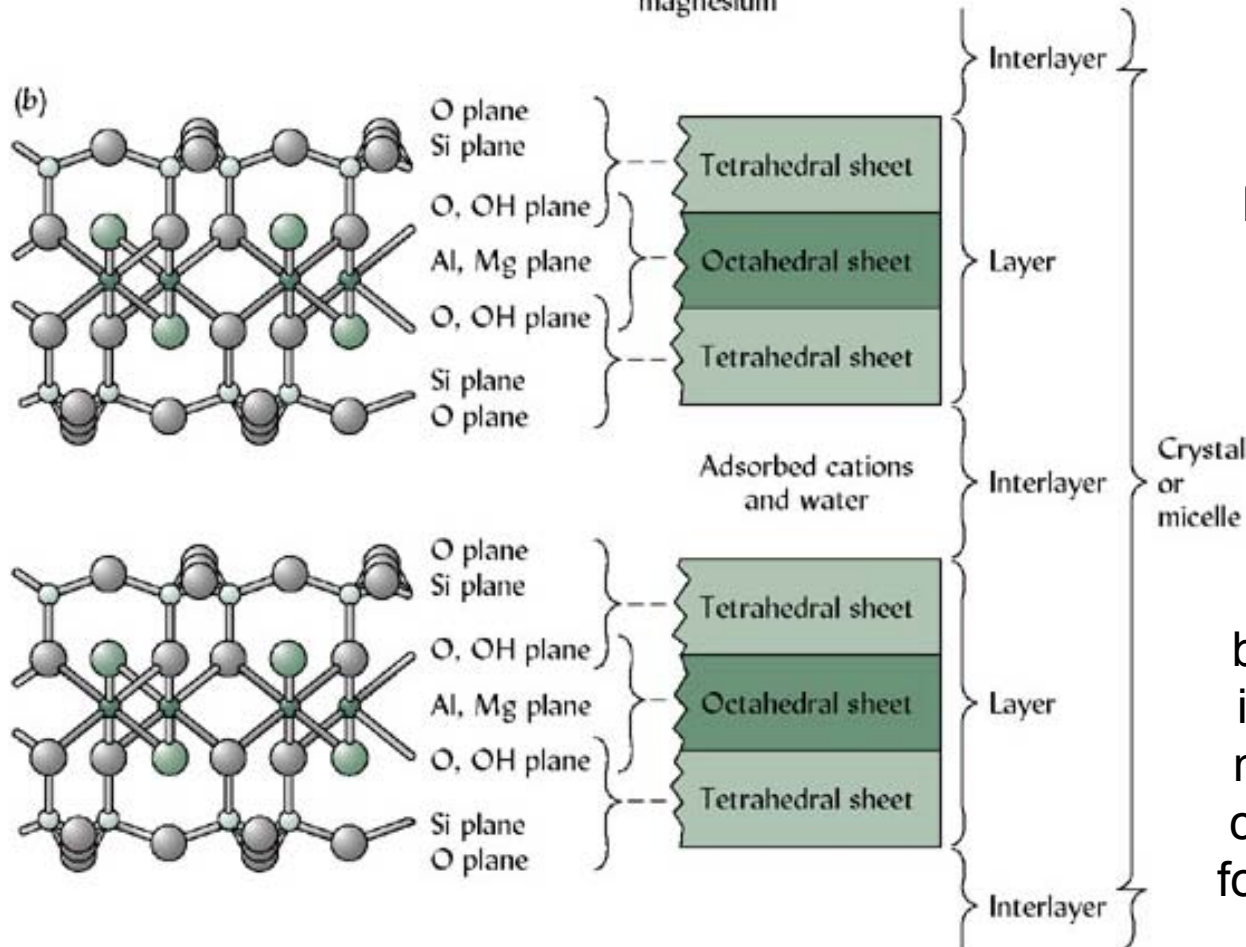


# Isomorphous Substitution

- In the tetrahedra,  $\text{Al}^{3+}$  ions substitute for  $\text{Si}^{4+}$  ions
- In the octahedra,  $\text{Fe}^{2+}$ ,  $\text{Mg}^{2+}$  ions substitute for  $\text{Al}^{3+}$  or  $\text{Fe}^{3+}$
- This leaves some of the negative charge unbalanced
- $\text{K}^+$ ,  $\text{Na}^+$ ,  $\text{Ca}^{2+}$ ,  $\text{Mg}^{2+}$  and other ions can be absorbed on the edges and between sheets to restore charge balance



The tetrahedral and octahedral sheets are the fundamental structural units of silicate clays



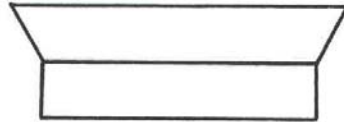
2 to 4 sheets can be stacked in sandwich-like arrangements (called layers) bound together by the sharing of oxygen atoms

The interlayer zones, binding the layers together is largely controlled by the nature of atom and charge combinations of the sheets forming the individual layers

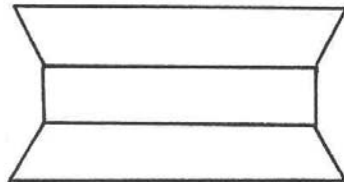
## Silicate Clays

### Main groups:

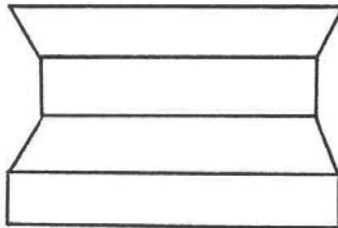
**1-to-1   silica sheet + alumina sheet = layer**  
**(tetrahedral)   (octahedral)**



**2-to-1   silica + alumina + silica = layer**  
**(Mg, Fe)**



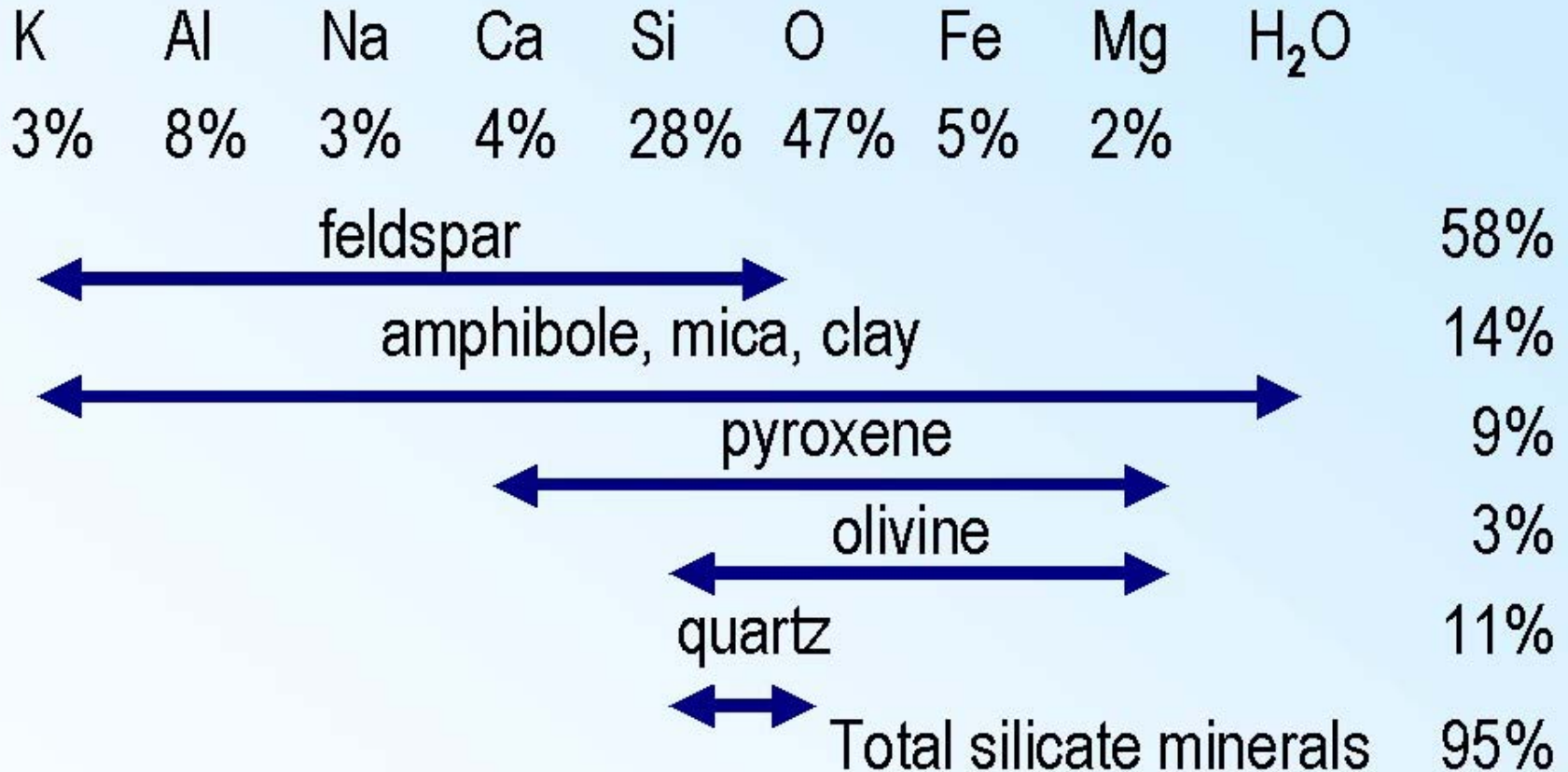
**2-to-1-to-1**  
**silica + alumina + silica + magnesia = layer**



**All are based on the layer (sheet) silicate**

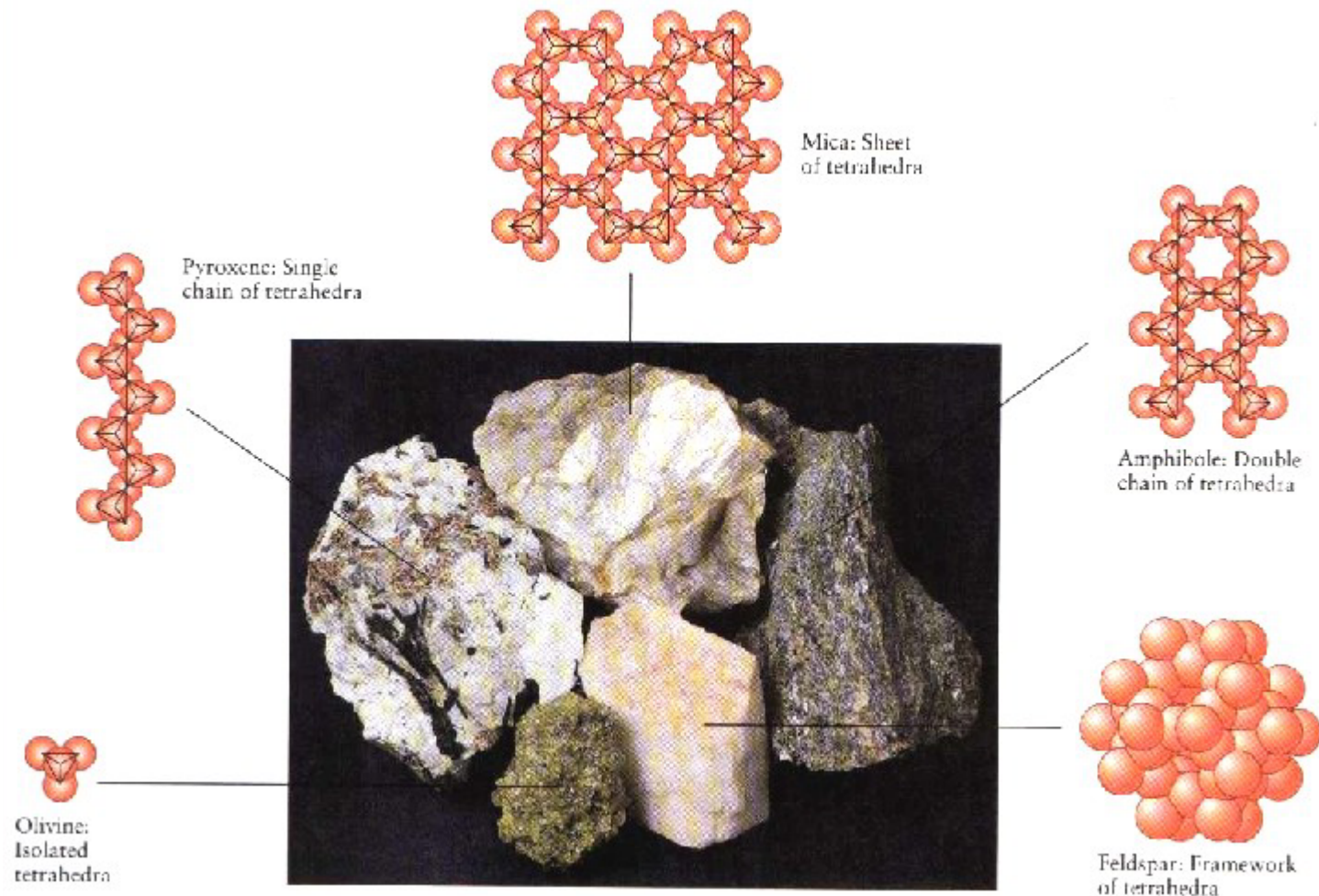
Chemical Weathering: Process that changes minerals from their original composition to a new composition

## Silicate Minerals in Earth's Crust



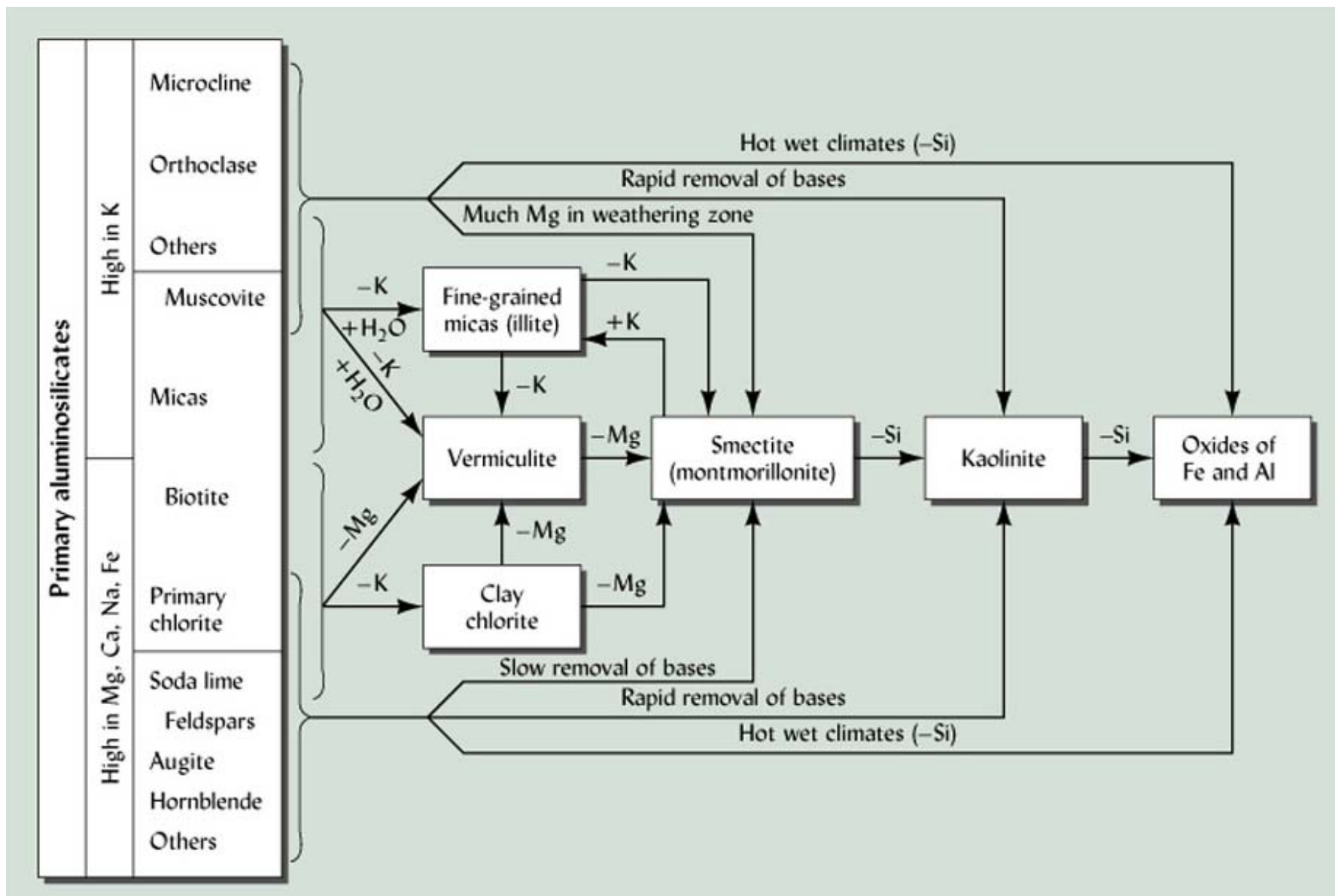


# Linked $\text{SiO}_4$ tetrahedra form structural backbone for silicate mineral groups

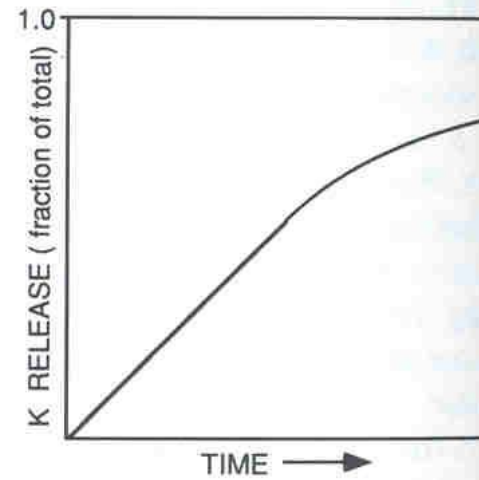
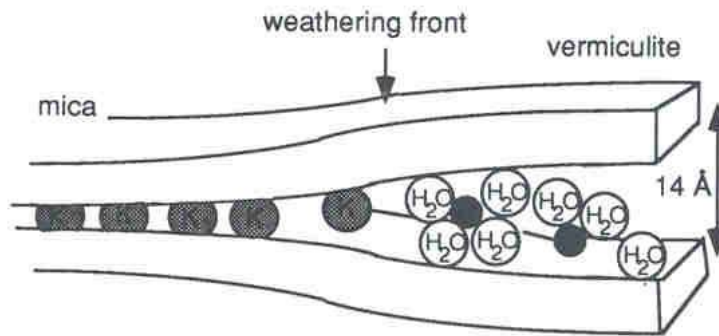


Secondary phyllosilicate mineral formation is not always from scratch, but is generally in a dynamic interaction with chemical weathering – ie transformation

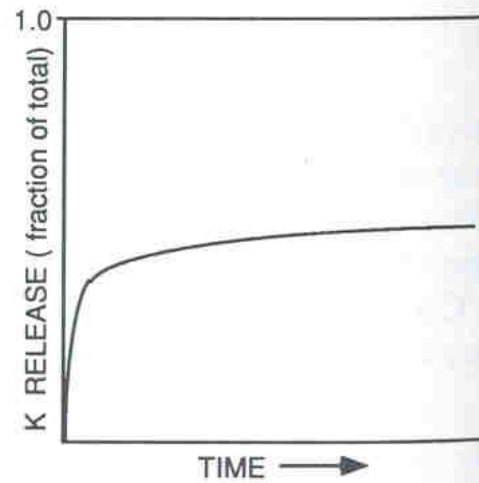
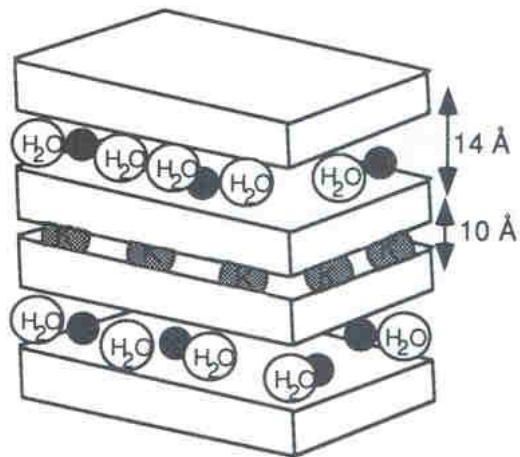
### Alteration versus Recrystallization



### EDGE WEATHERING

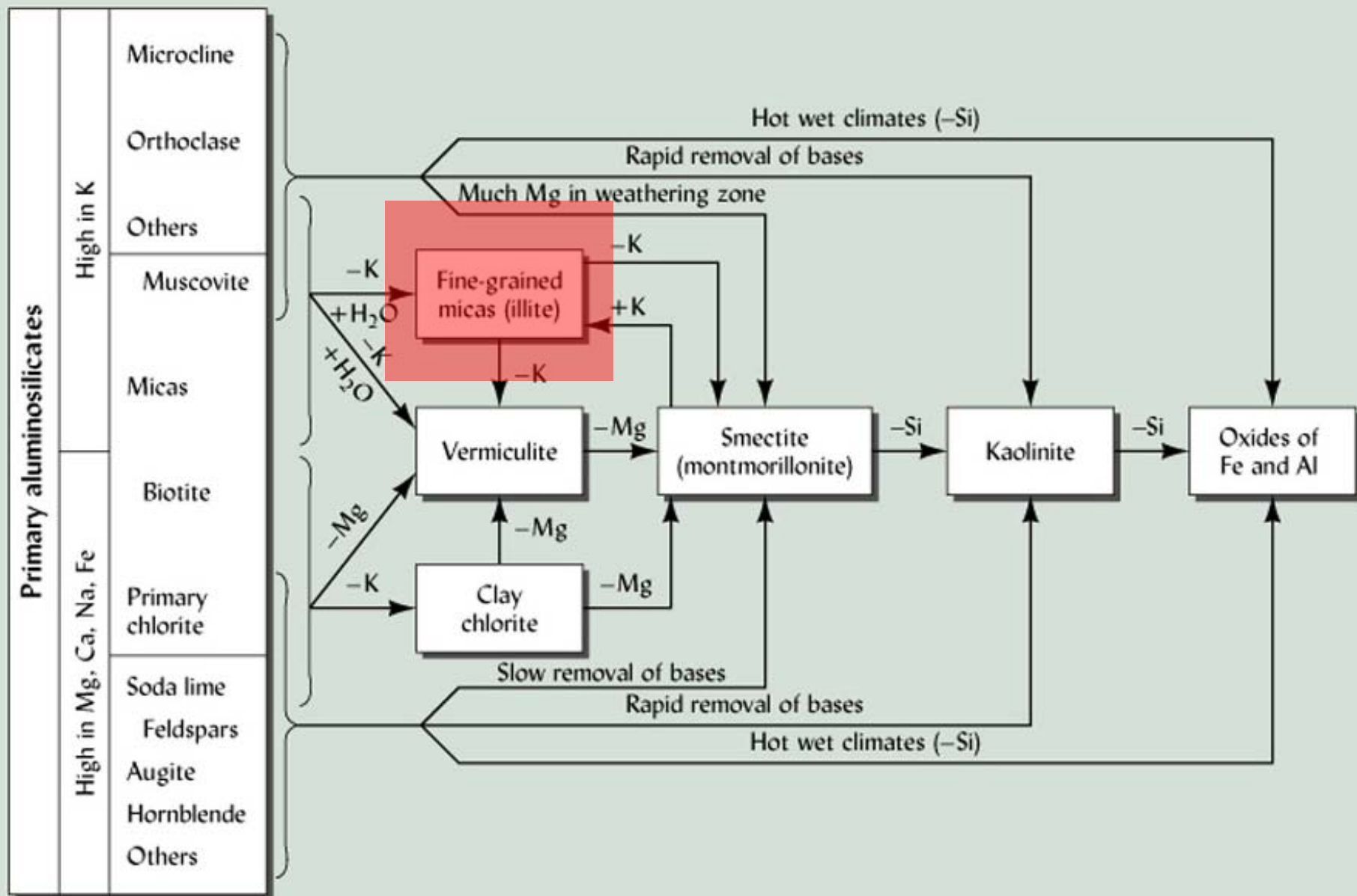


### LAYER WEATHERING



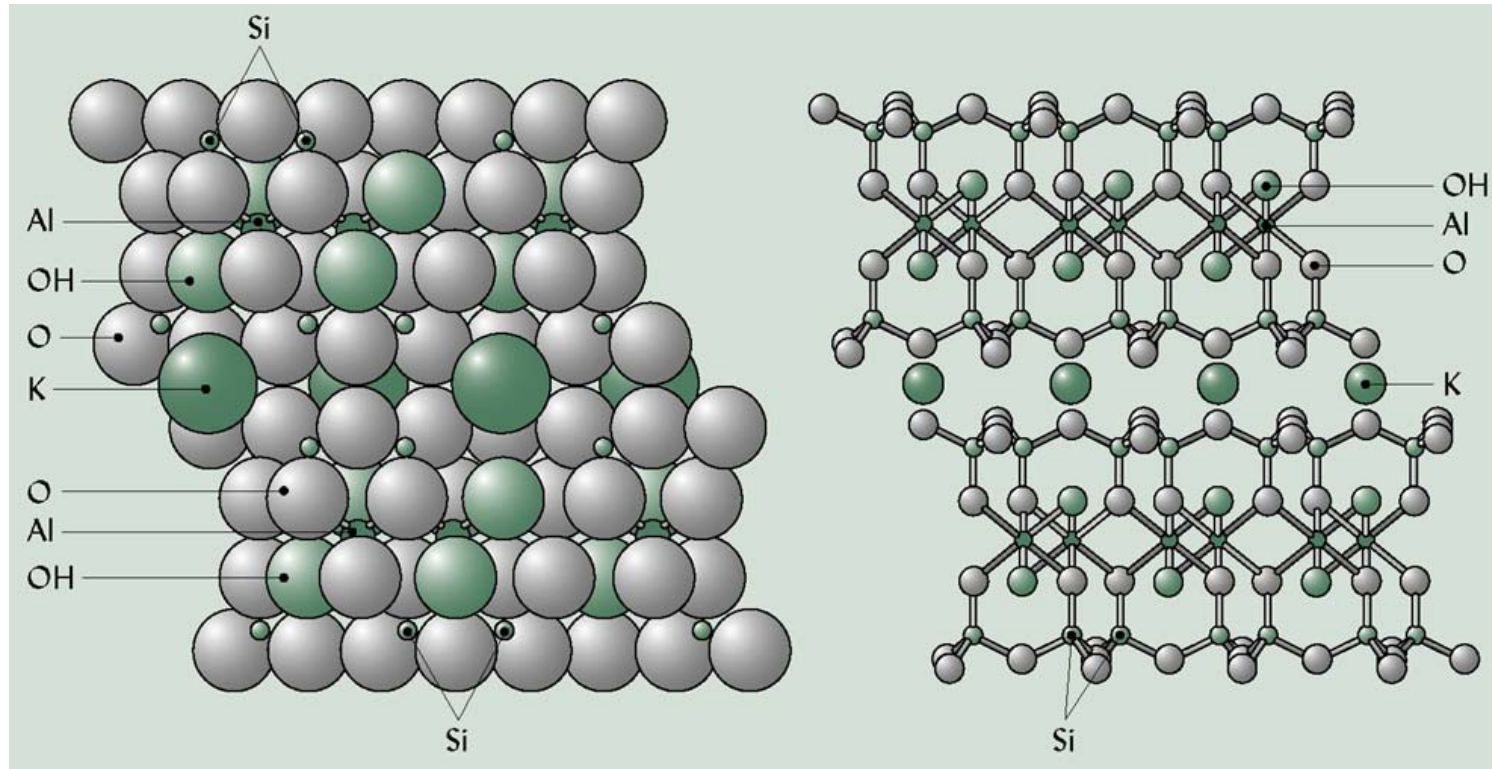
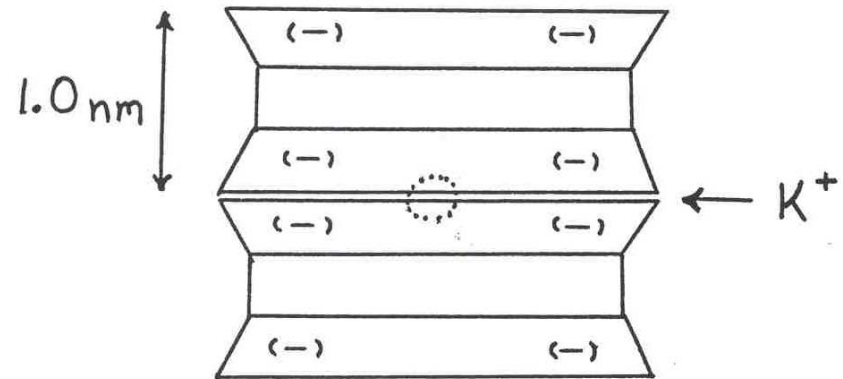
**Figure 6.4.** Schematic picture of edge weathering of large mica particles and layer weathering of small mica particles.

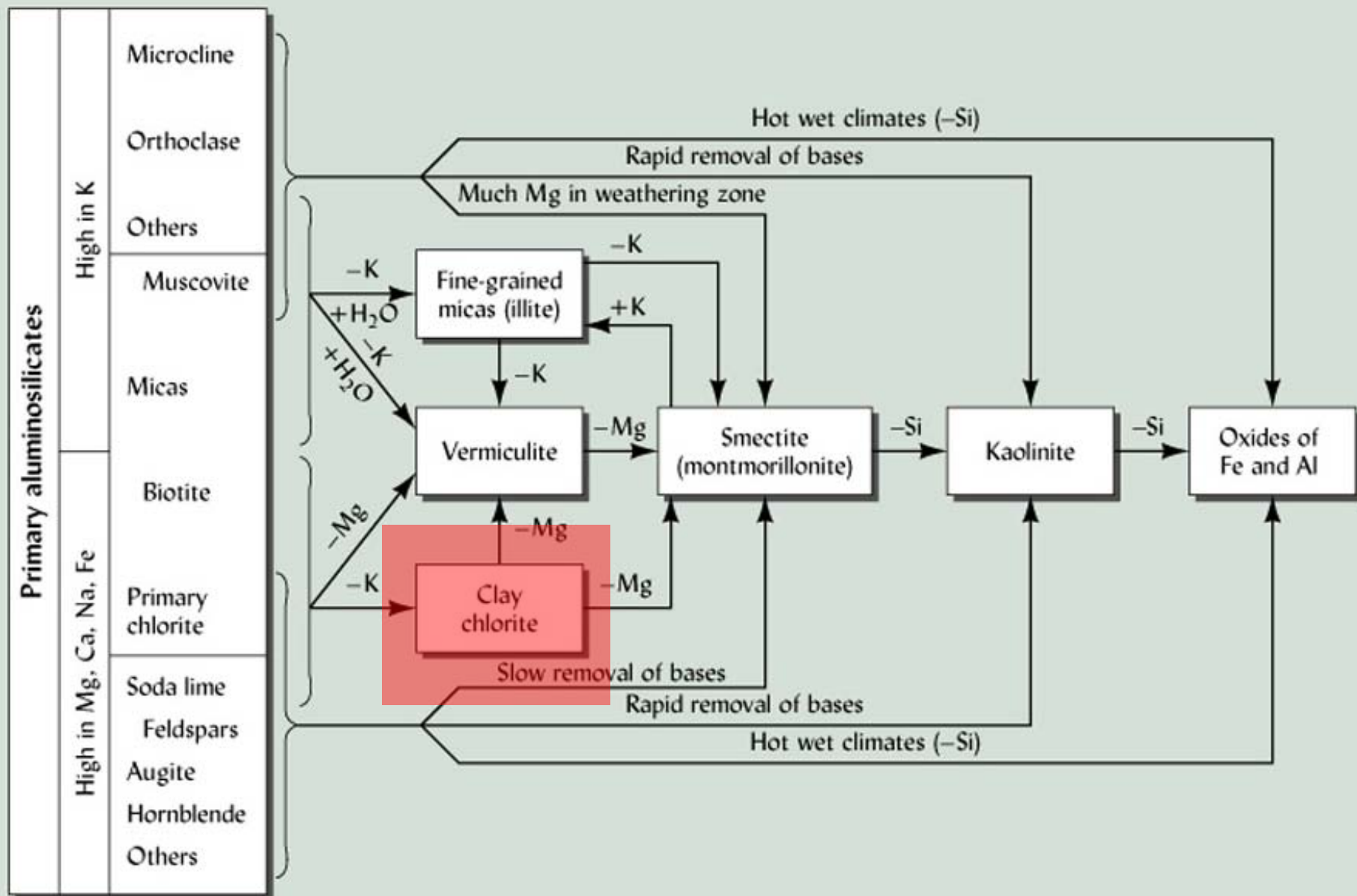




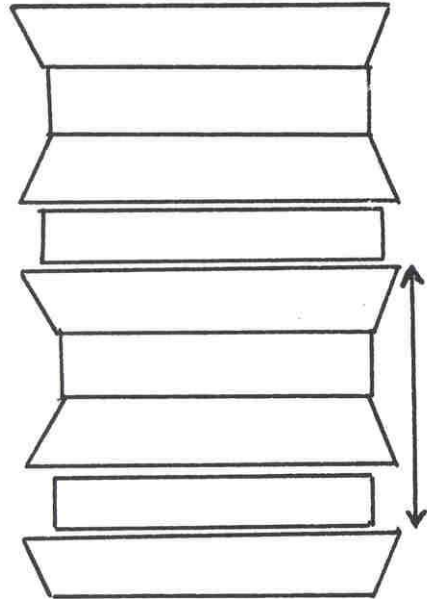
# Illite

- **2-to-1, non-expanding clay mineral**
- **often called "hydrous mica"**
- **interlayer  $K^+$  (as in mica) prevents swelling**



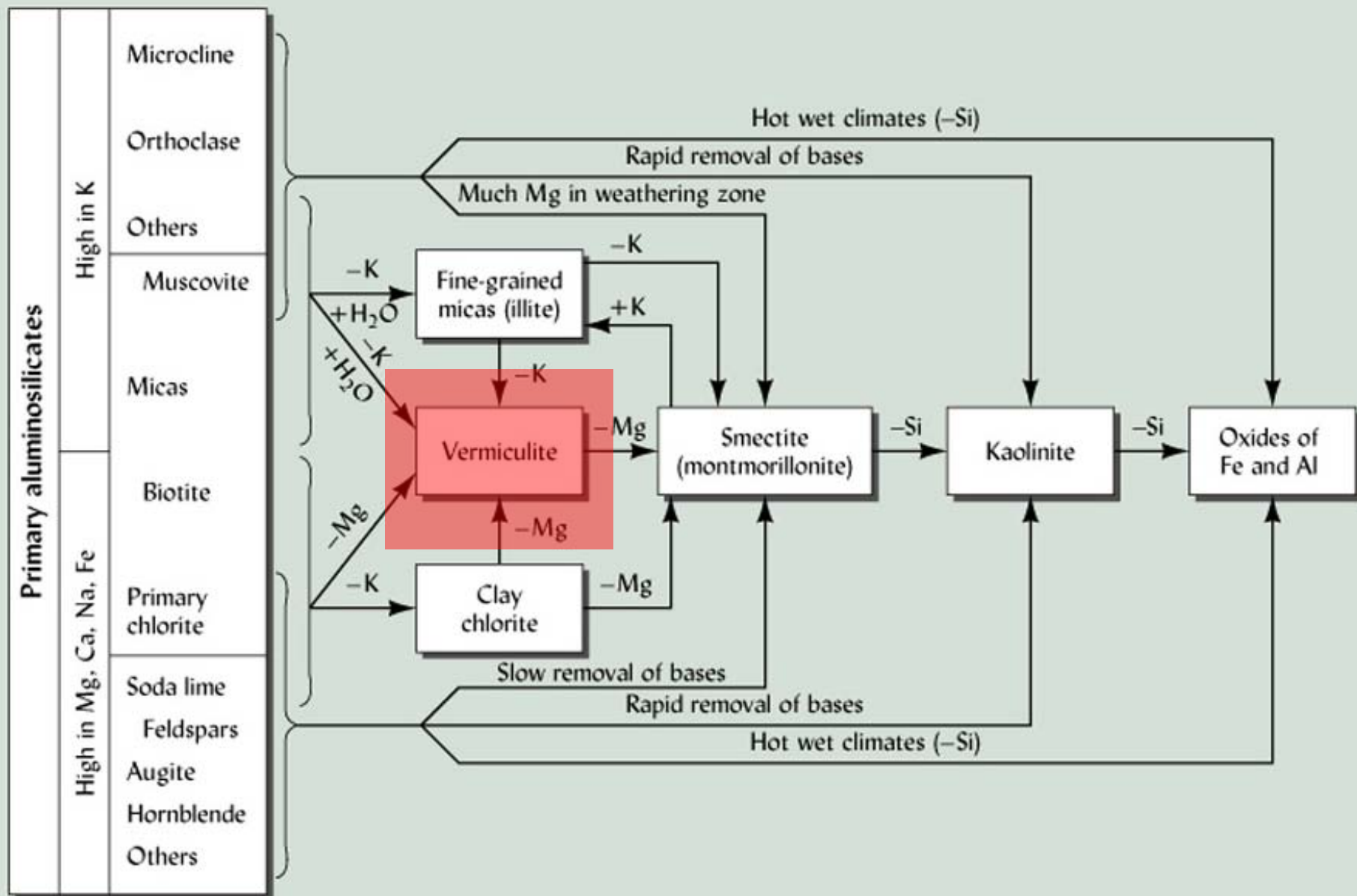


## Chlorite:



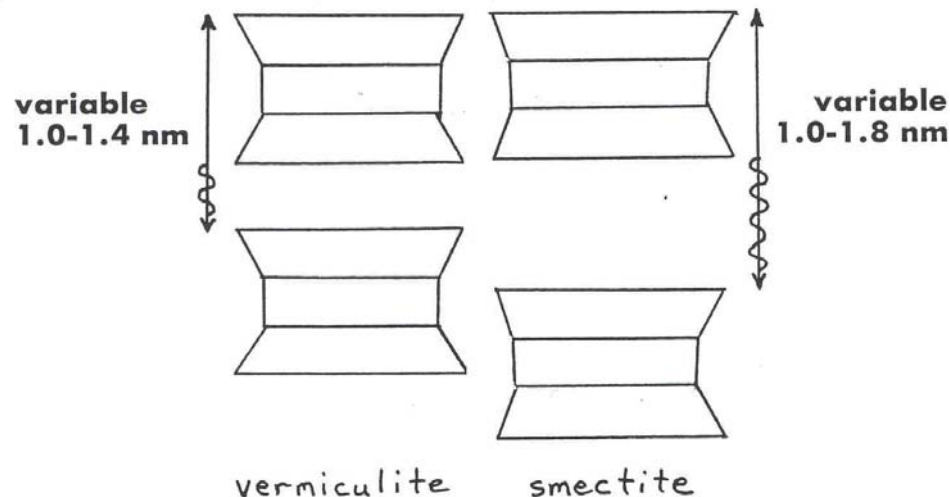
**2-to-1-to-1, non-expanding clay mineral like vermiculite, but has positively charged Mg hydroxide sheet replacing exchange cations between 2-to-1 sheets:  $[\text{Mg}_5\text{Al}(\text{OH})_{12}^+]$**

**1.4 nm**

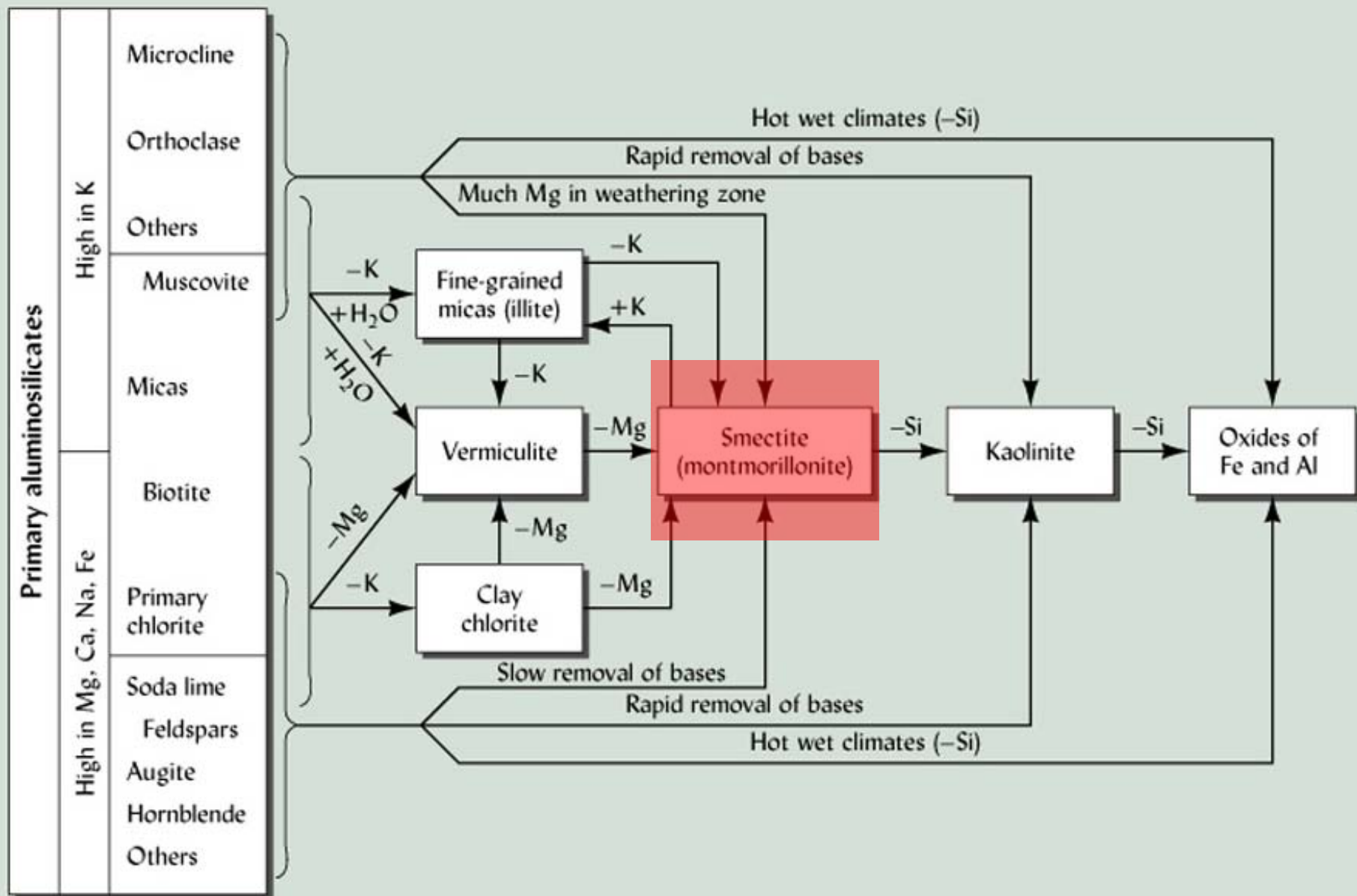


### **Vermiculite:**

- **important 2-to-1 clay mineral in soils**
- **structurally similar to smectites with 3 differences:**
  1. **most or all of isomorphous substitution is tetrahedral**
  2. **isomorphous substitution is greater**
    - $\text{Al}^{3+}$  for  $\text{Si}^{4+}$  (tet.)
    - $\text{Fe}^{3+}$ ,  $\text{Fe}^{2+}$ ,  $\text{Mg}^{2+}$  for  $\text{Al}^{3+}$  (oct.)
  3.  $\text{Mg}^{2+}$  is the usual exchangeable cation
- **result of these differences:**
  1. **swelling in water limited to 1.4 nm**
  2. **larger capacity for adsorption of exchangeable cations than smectites**



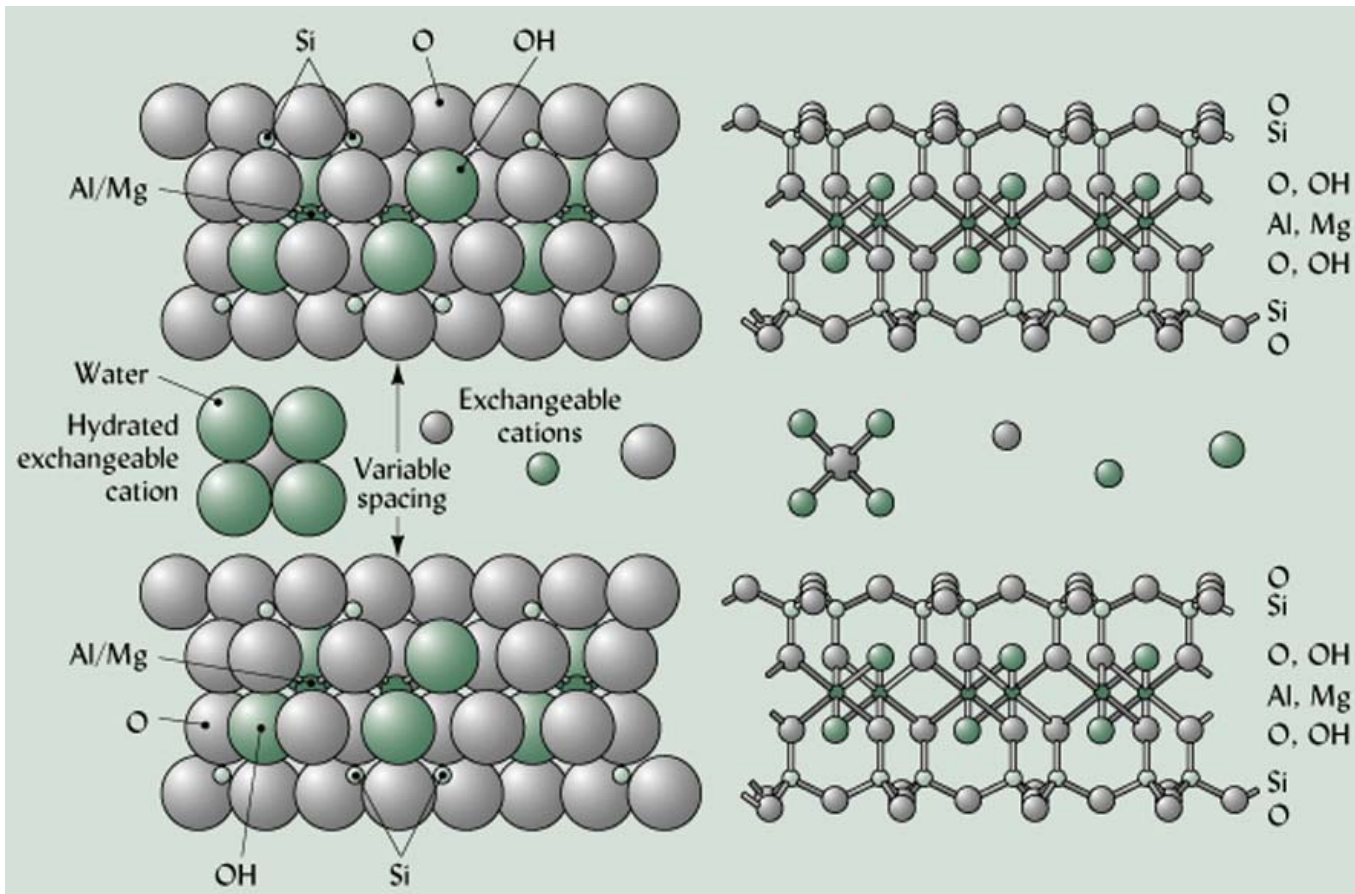




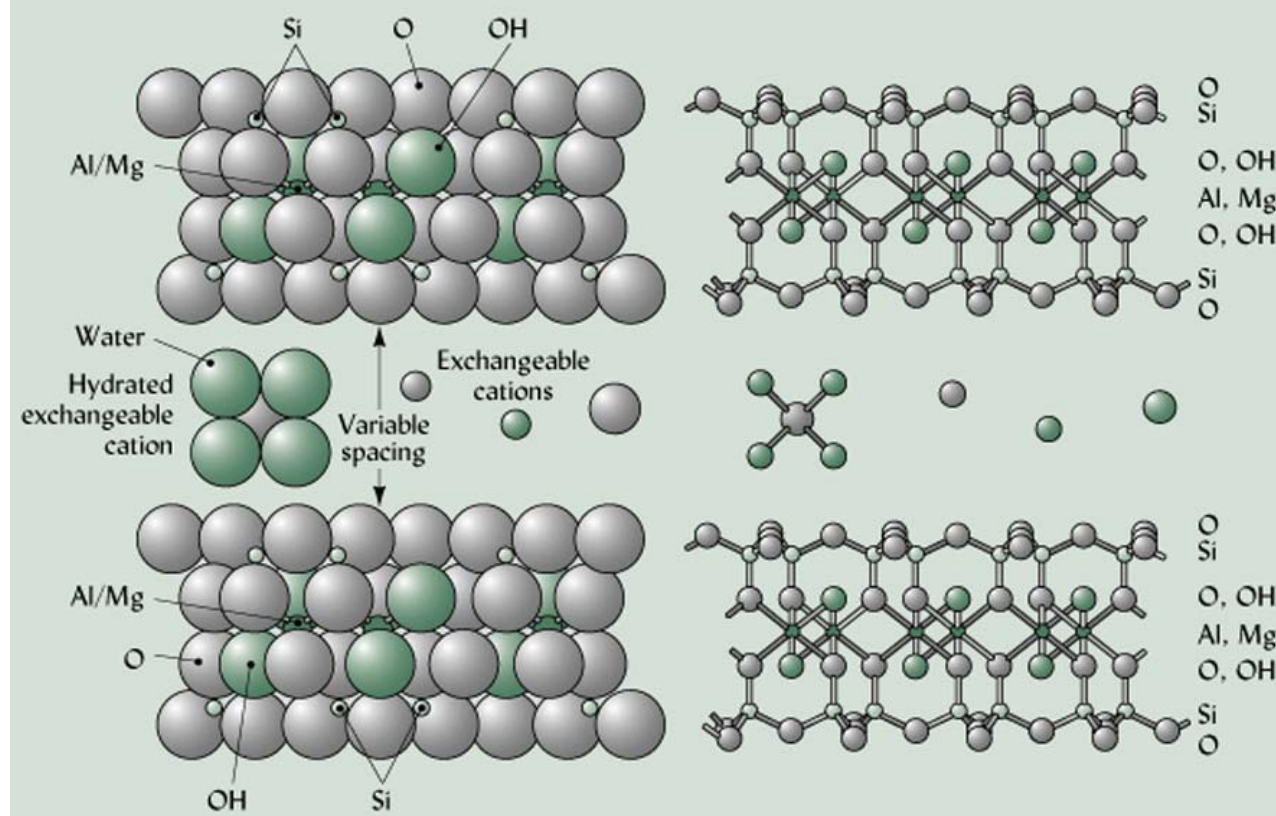
## Smectites (a family of minerals):

- **important 2-to-1 clay minerals in soils**
- **each layer has octahedral sheet sandwiched between two tetrahedral sheets**
- **formula of unit cell is variable, but many common smectites are based on pyrophyllite:**

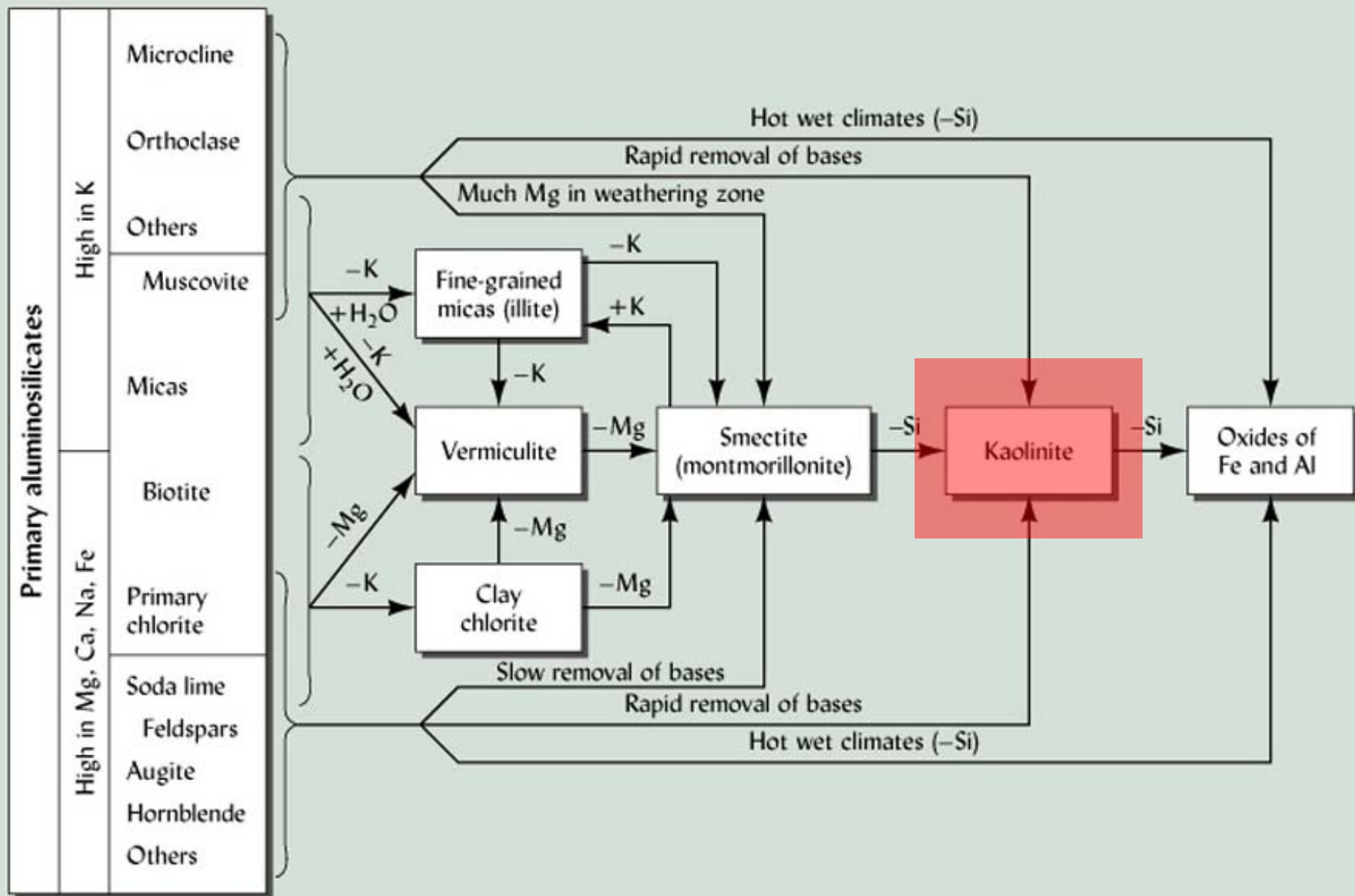
sheets held together at apical oxygen ions







- **able to expand in water, increasing size of unit cell & amount of adsorbed water**
- **shrink/swell behavior is related to the loss or gain of water in the interlayer space**
- **adsorbed water ( $n \text{ H}_2\text{O}$ ) is associated with exchangeable cations**
- **huge surface area due to exposed interlayer surface ( $800 \text{ m}^2/\text{g}$ )**

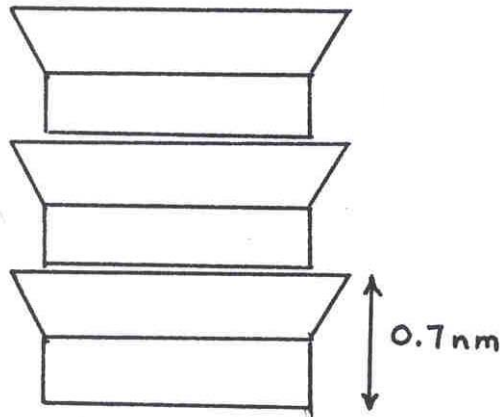


## Kaolinite:

- most important 1-to-1 mineral in soil
- formula shows 1:1 silica/alumina ratio



**UNIT CELL =** minimum unit which can be repeated to form the crystal structure



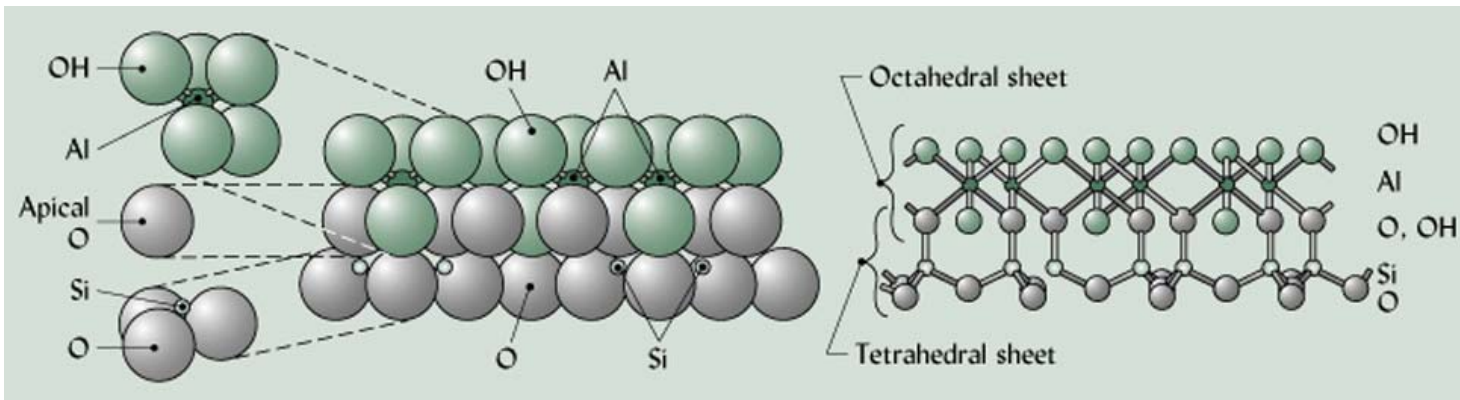
each layer (unit cell) is 0.7 nm thick

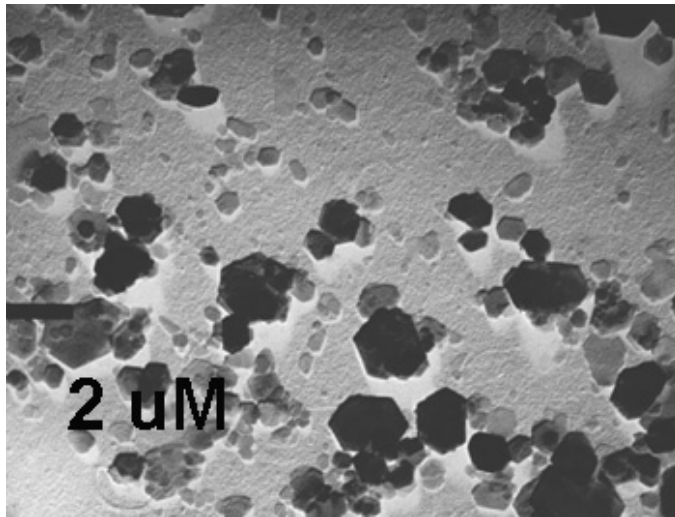
takes 3,000 layers to form a .002 mm thick clay particle

hydrogen bonds hold layers together (hydroxyl contacts oxygen on adjacent layer)

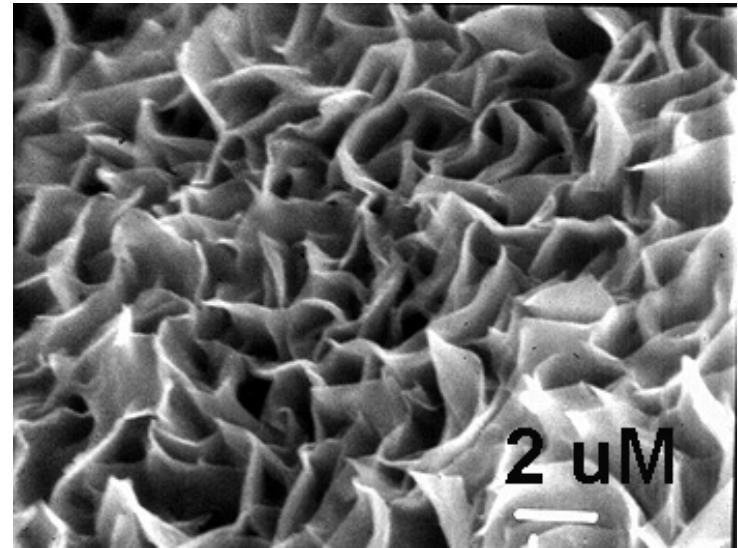
non-expanding in water

little ability to retain cations





1: 1 kaolinite



2:1 smectite

